



# ESA-JAXA Pre-Launch EarthCARE Science and Validation Workshop

13 – 17 November 2023 | ESA-ESRIN, Frascati (Rome), Italy

## Development of a new simulator on COSP2 for vertical doppler velocity of EarthCARE CPR

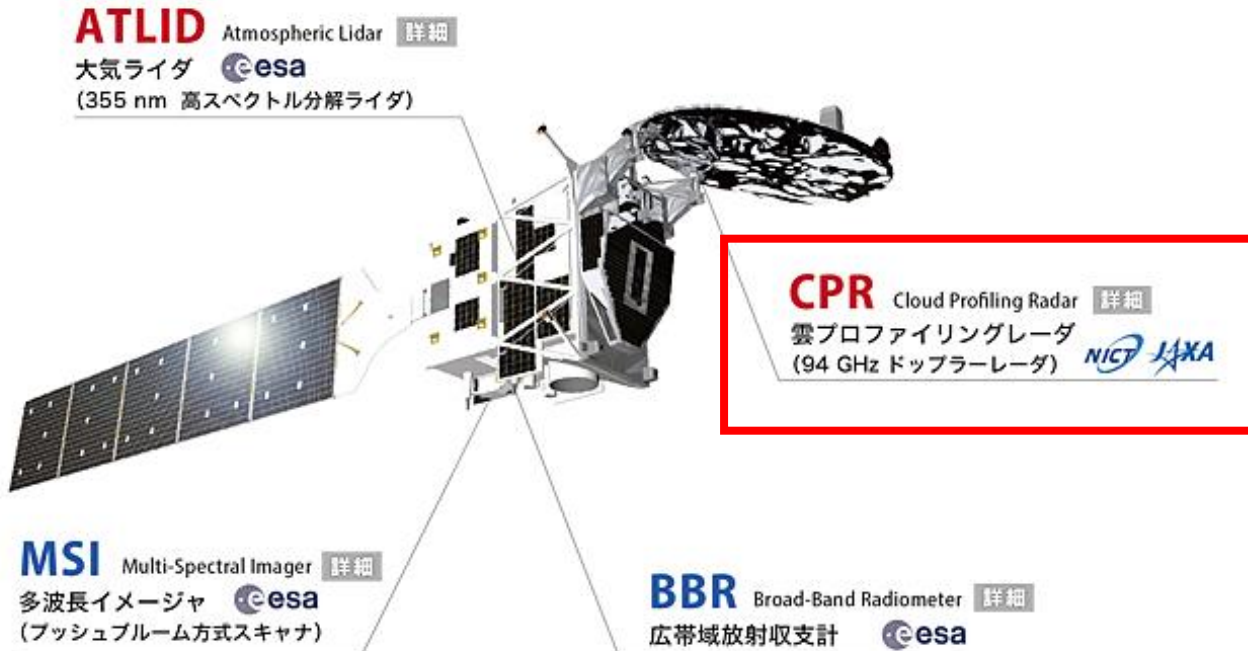
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*2: National Institute of Information and Communications Technology, Japan*

**See also Poster #11**

# Newer observation by EarthCARE CPR and satellite simulator on COSP2



Picture from JAXA webpage (14 Nov. 2022 browsed)

[https://www.eorc.jaxa.jp/EARTHCARE/about/hardware\\_j.html](https://www.eorc.jaxa.jp/EARTHCARE/about/hardware_j.html)

- Higher sensitivity than CloudSat
- **Doppler velocity observation**  
→ Dynamics of cloud microphysics

$$v_d = w + \frac{\int_{r_{\min}}^{r_{\max}} \underbrace{dn(r)/dr}_{\text{size distribution}} \cdot \underbrace{C_{\text{bk}}(r)}_{\text{back scatter}} \cdot \underbrace{v_f(r)}_{\text{fall velocity}} dr}{\int_{r_{\min}}^{r_{\max}} \underbrace{dn(r)/dr \cdot C_{\text{bk}}(r)}_{\text{radar reflectivity factor}} dr}$$

Labels in the diagram:  
 - size distribution (green box)  
 - back scatter (blue box)  
 - fall velocity (red box)  
 - vertical air motion (orange box)  
 - radar reflectivity factor (blue box)

- Retrieval formulation
- What is needed for the new simulator...  
→ droplet fall velocity  
→ vertical air motion
- For size distribution and back scatter, use codes already implemented in COSP2.

# Contributions to climate models expected from new doppler velocity observation

- Droplet fall velocity was often considered as a tuning parameter. It affected the performance of models, both GCMs and CRMs.  
→ Risk of a physically meaningless setting.
- ✓ New observation is expected to constraint the fall velocity of microphysics.
- Observations of cumulus mass flux itself were very limited.
- ✓ Quantitative estimation will be provided by EarthCARE doppler velocity.  
→ But CPR attenuation would be a difficult issue...

To maximize these benefits from EarthCARE,  
it is necessary to ensure a basis for comparing models with observation.

- 1. To develop a new simulator for doppler velocity on COSP2.**
- 2. To compare ground-based doppler radar and GCM.**
- 3. Further discussion of an impact for climate from droplet fall velocity and use of doppler velocity statistics.**

# Ground-based radar

## 1. NICT

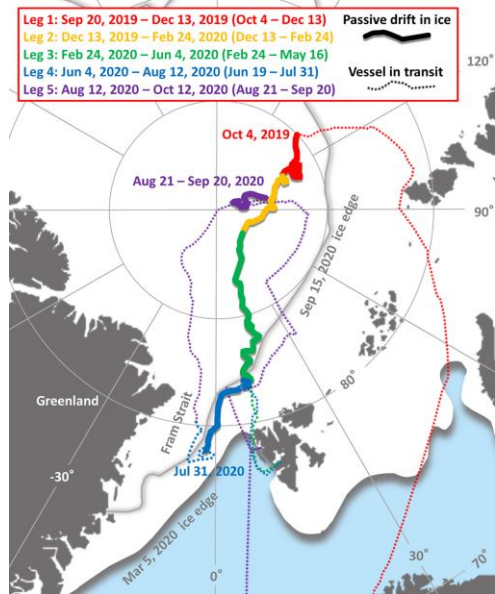
(National Institute of Information and Communications Technology)  
provided by Horie-san (NICT).  
In Koganei City, Tokyo. (30km west of Senso-ji temple, Asakusa)



## 2. MOSAIC

(Multidisciplinary drifting Observatory for the Study of Arctic Climate)  
details: <https://mosaic-expedition.org>  
provided by ARM, ship-borne  
drifting on sea ice in Arctic Sea

Both datasets are  
in Aug. 2020.



# GCM setup

## MIROC6 (Tatebe et al. 2019; GMD)

- Prognostic precipitation scheme  
(Michibata et al. 2019; JAMES)

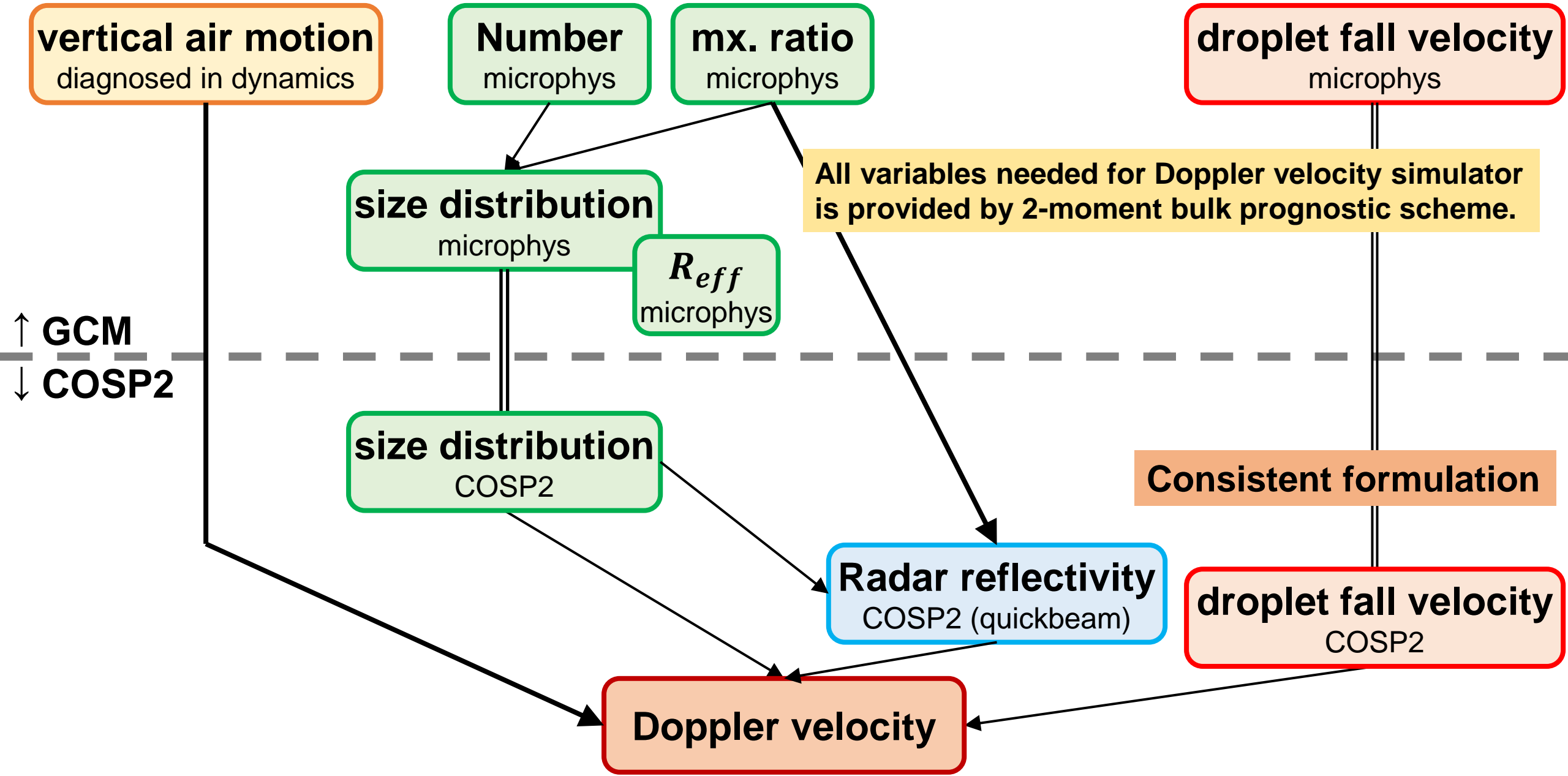
This bulk scheme prognoses mass and number of falling hydrometeors: cloud, rain, and snow.

- about 100 km resolution (t85 l40h)
- **JJA 2020**
  - ✓ 5×5 grids centered on Tokyo (NICT)
  - ✓ 30W-20E, 75N-88N (MOSAIC)

## COSP2

- CFMIP Observation Simulator Package
- Doppler velocity simulator is developed.
- #subcolumn = 140

# Simulator design: prognostic precipitation scheme



# Construction of droplet fall velocity $v_f$

$$v_f = \text{viscous drag} \times \text{formulation}$$

- **viscous drag**

The higher the air density, the greater the viscous drag and the slower the fall speed.

$\sqrt{\rho_0/\rho}$ : Square root of the ratio of air density  $\rho$  to the standard value ( $\rho_0$ ; 1013 hPa, 0°C)  
on/off switchable

- **formulation**

2 types of functions of droplet diameter  $D$  for  $v_f$  formulation

1. **Power law:  $aD^b$**

2. **PL08:  $b_1 - b_2 \exp(-b_3 D) + (b_2 - b_1) \exp(-5b_3 D)$**

Posselt and Lohmann (2008; ACP)

ctrl	large-scale condensation					cumulus convection				
	droplet	cloud liq.	cloud ice	rain	snow	graupel	cloud liq.	cloud ice	rain	snow
viscous drag	×	×	×	×	×	×	×	×	○	○
PL08	○	×	○	×	×	×	○	×	×	×
Power law; $a$	-	1.107	-	3.321	19.3	-	1.107	842	4.84	
Power law; $b$	-	0.22	-	0.22	0.37	-	0.22	0.8	0.25	

**PL08:  $b_1 = 9.65, b_2 = 10.43, b_3 = 600$**

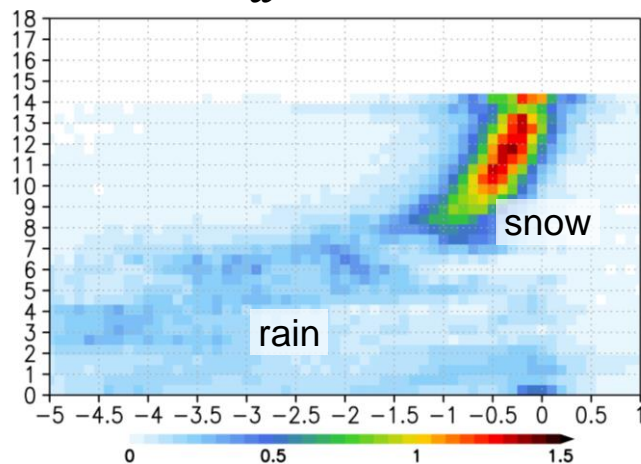
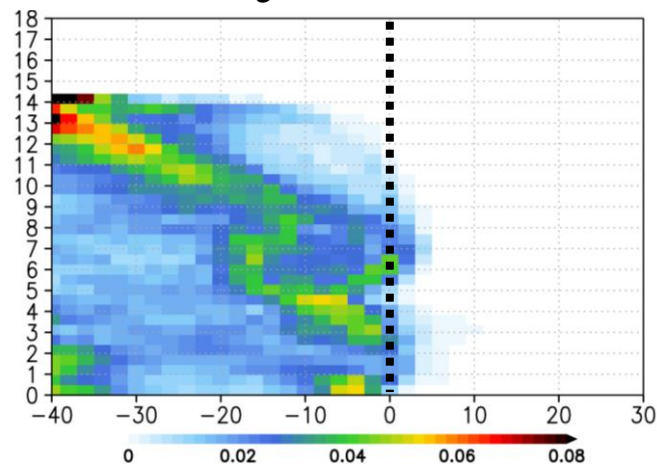
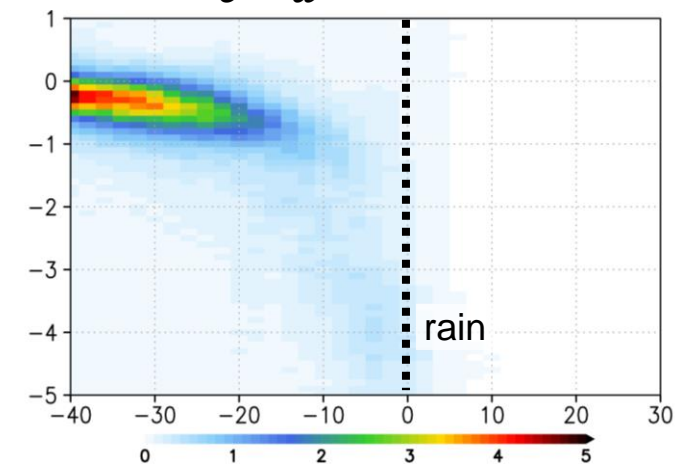
# Comparison between NICT obs. and MIROC6

2020/08

observation

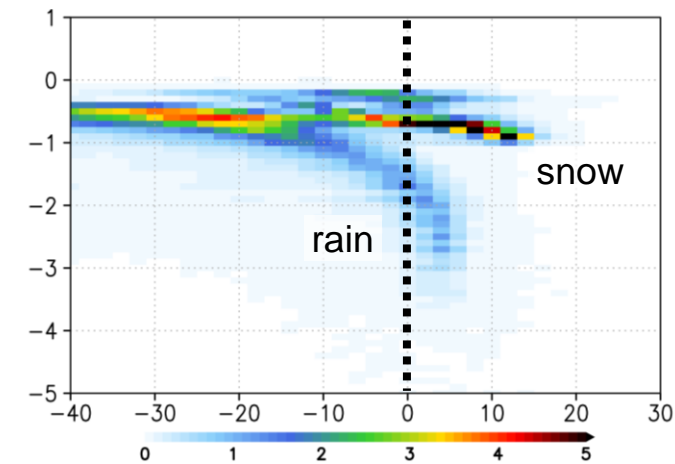
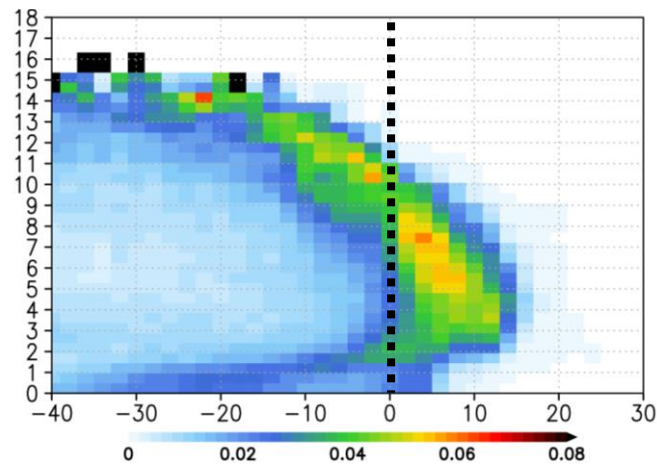
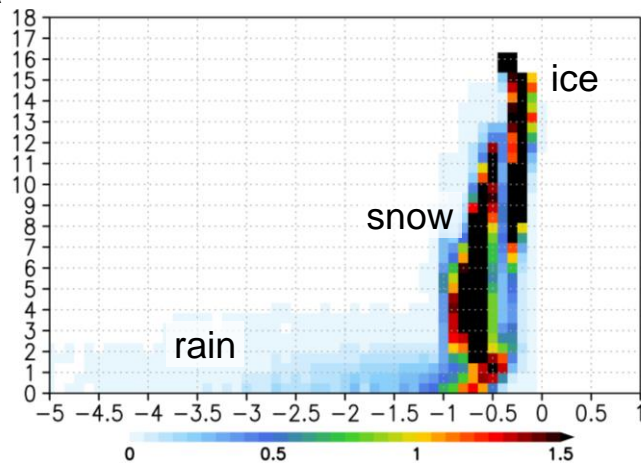
 $v_d$  CFAD

NICT

 $Z_e$  CFAD $Z_e-v_d$  2dPDF

2020 JJA

MIROC6 ctrl



- Slower  $v_d$  in MIROC6, especially rain particle
- Discrete distribution, ice-snow

- Overestimating  $Z_e$   
✓ Often seen bias

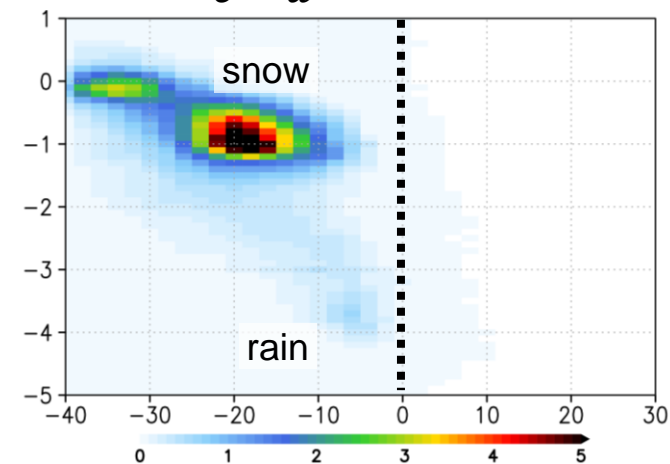
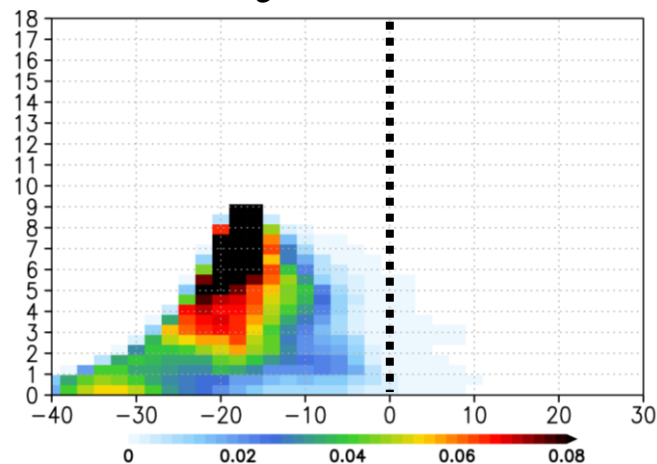
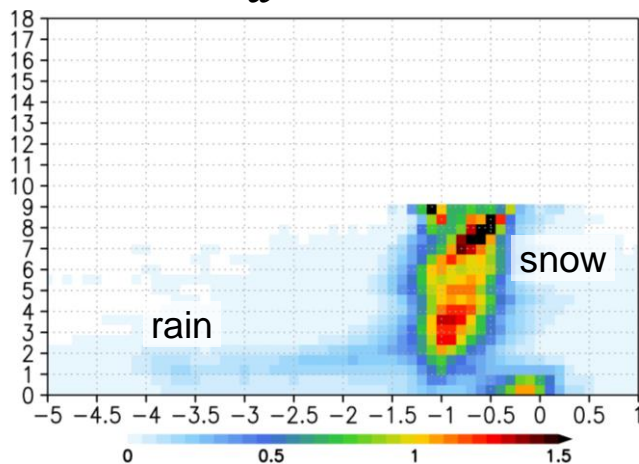
# Comparison between MOSAiC obs. and MIROC6

2020/08

observation

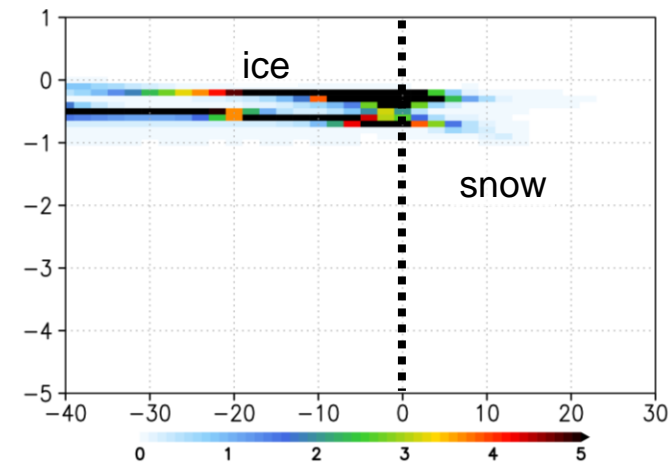
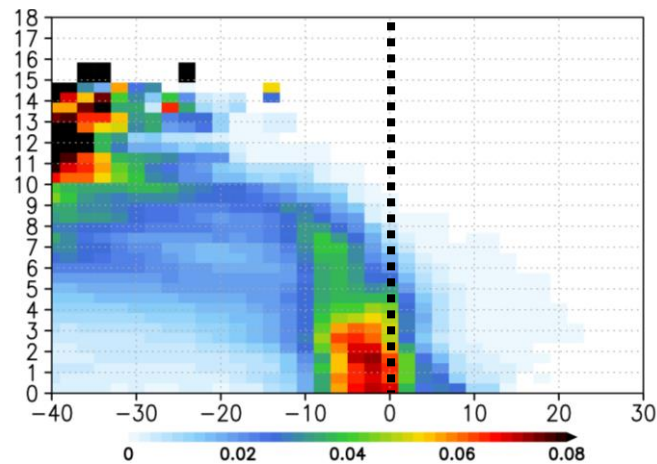
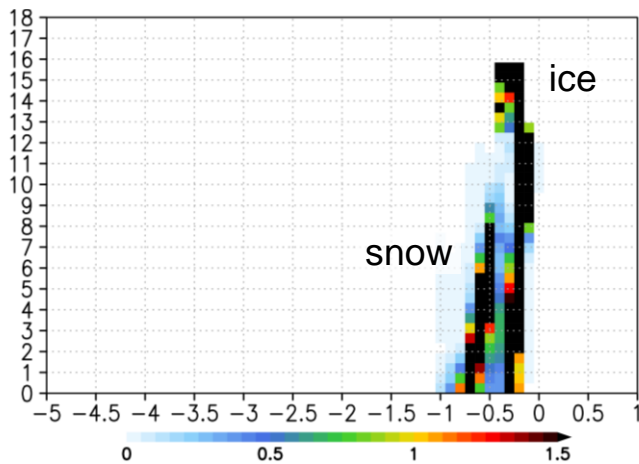
 $v_d$  CFAD $Z_e$  CFAD $Z_e - v_d$  2dPDF

MOSAiC



2020 JJA

MIROC ctrl



- Slower  $v_d$ , smaller variance in MIROC6.
- Lack of faster falling droplets below 2km
- Overestimating  $Z_e$
- Distribution is quite different.



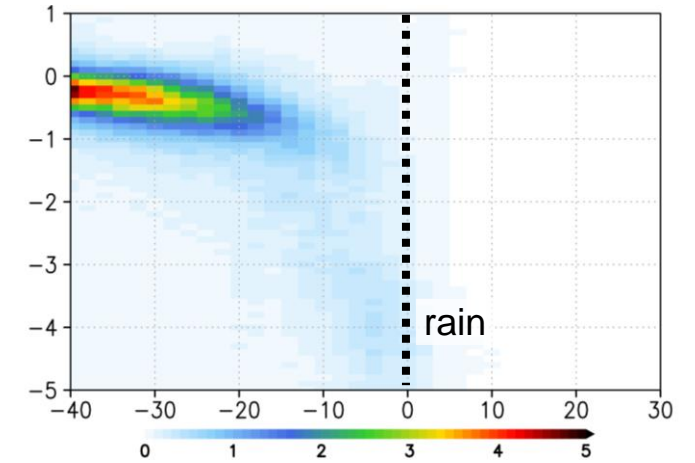
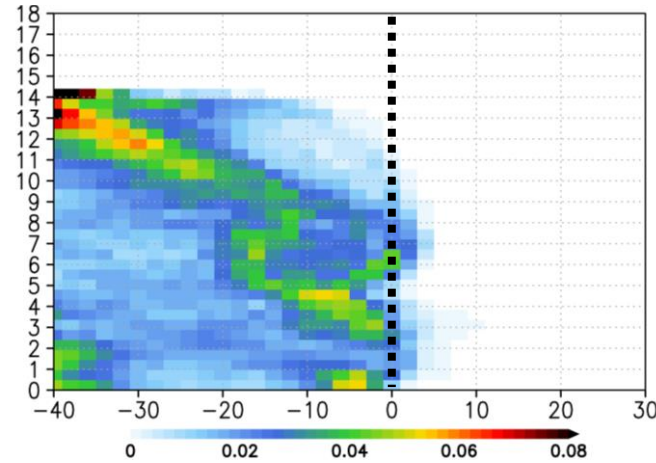
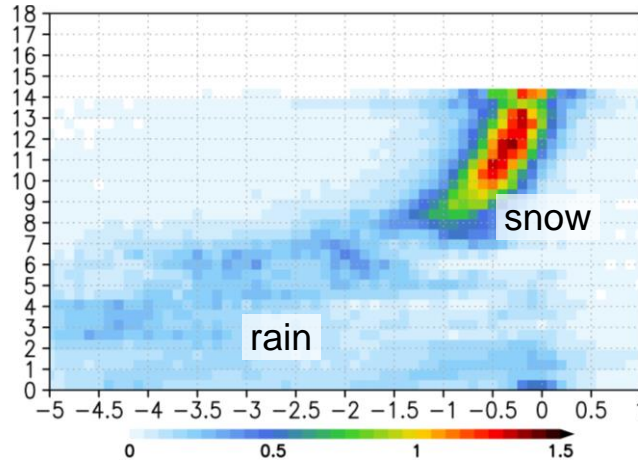
# Tuning to match $v_d$ in MIROC6 with NICT obs.

2020/08

observation

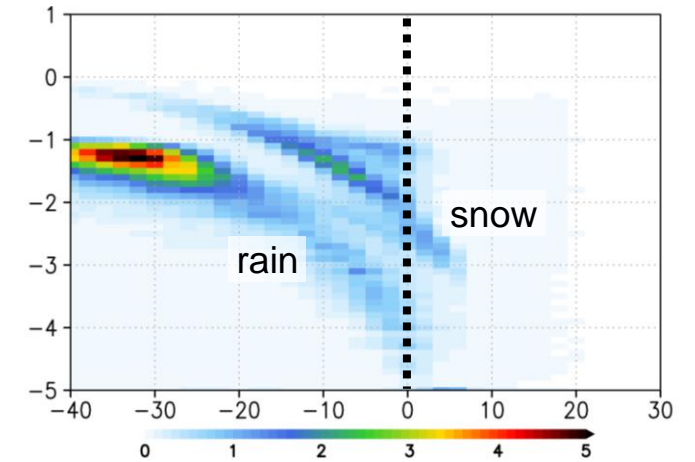
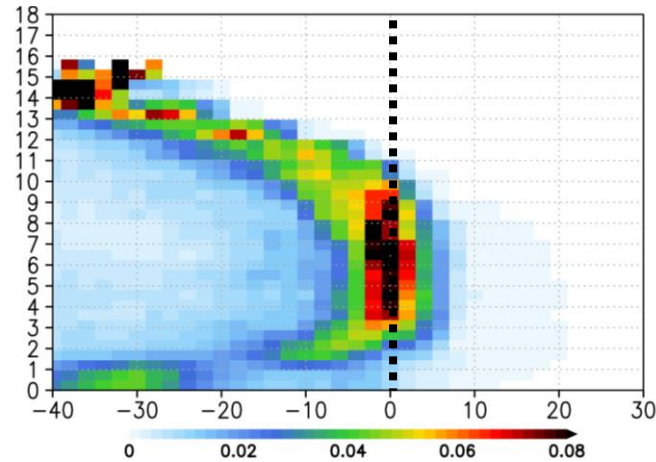
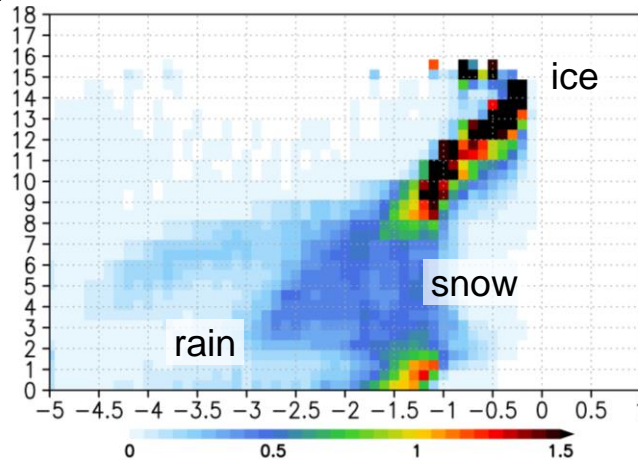
 $v_d$  CFAD $Z_e$  CFAD $Z_e$ - $v_d$  2dPDF

NICT



2020 JJA

MIROC ctrl

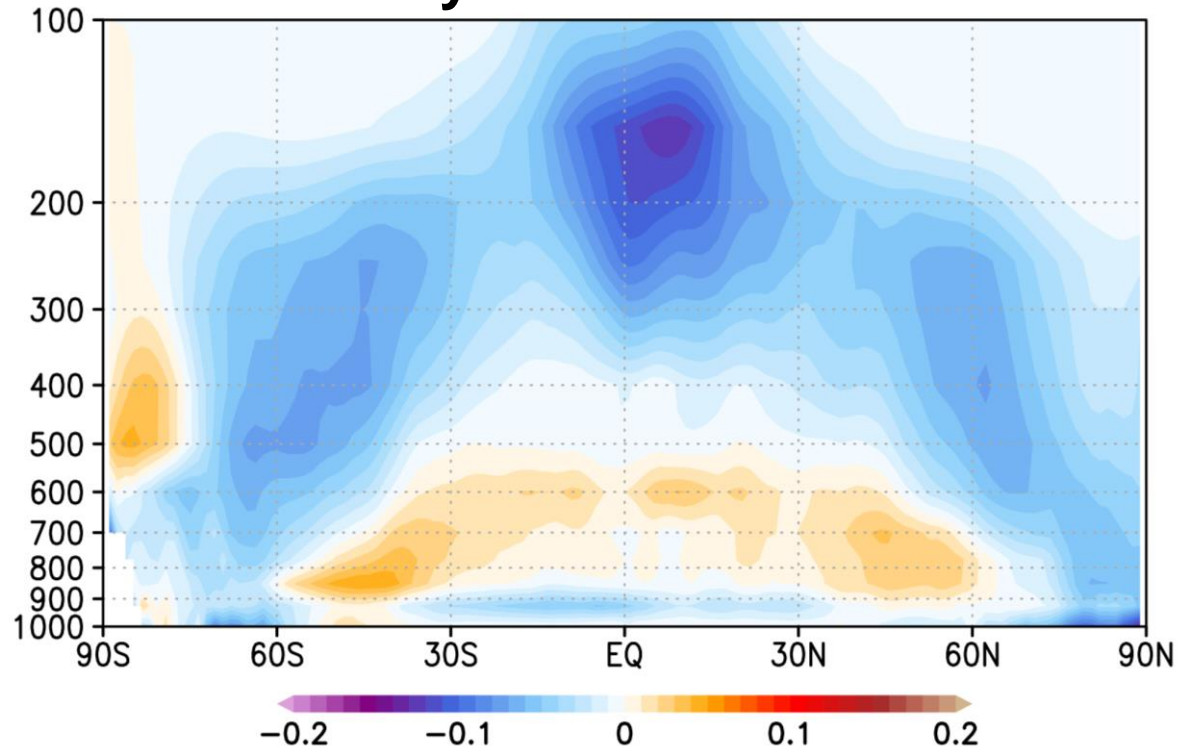


- Tuned to faster  $v_d$  and larger variance

- $Z_e$  becomes decreased
- better 2dPDF, but separated

# Impact on climate of $v_d$ tuning

## anomaly of cloud fraction



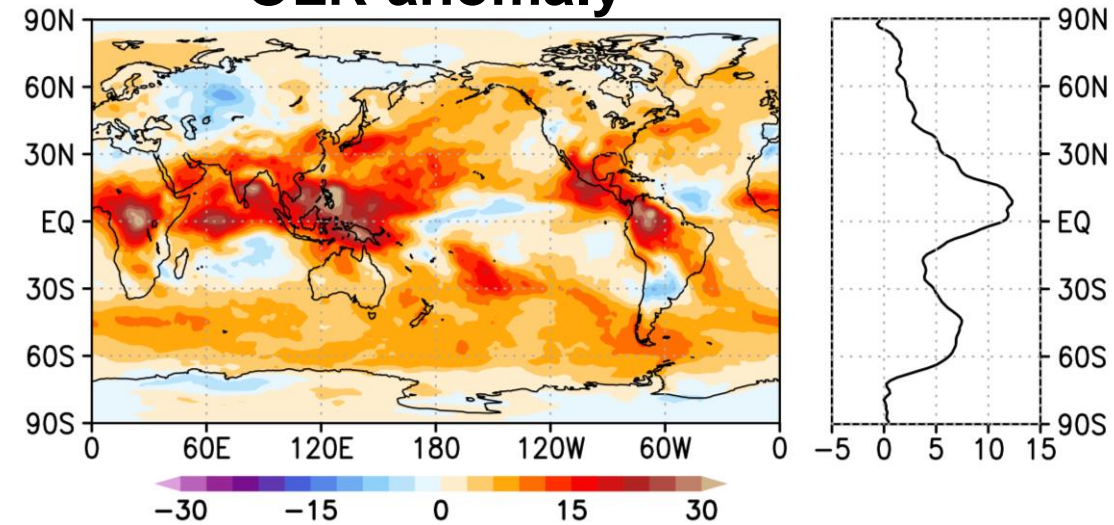
- Upper clouds strongly decreased.  
→ Due to faster fall and shorter lifetime.
- Lower clouds increased. → **Cooling**

## How to reduce this perturbation?

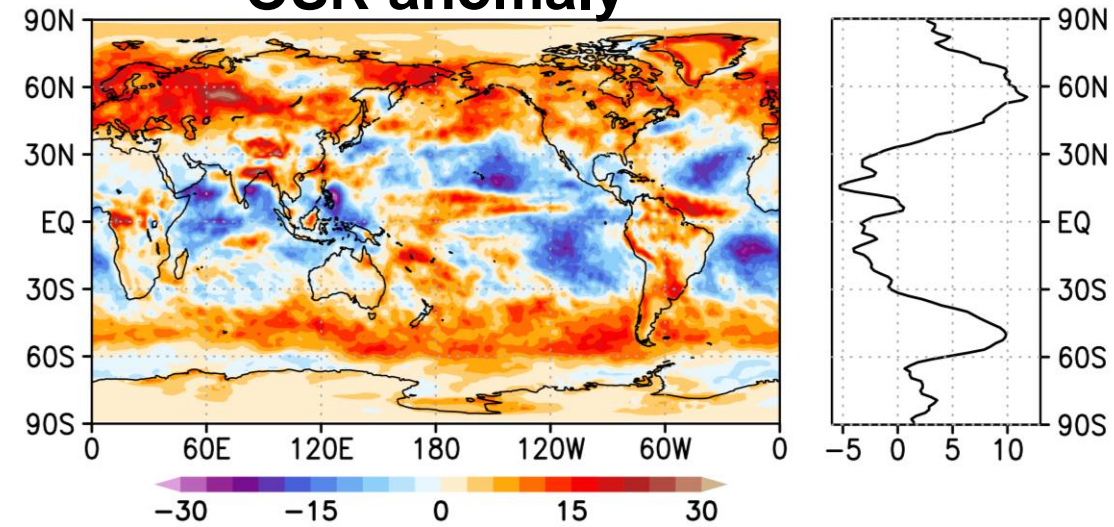
for liquid: autoconversion ratio, background CCN  
for ice: INP, shape, WBF process and mode...

5-year run from Jun. 2020  
Anomalies from control run are shown.

## OLR anomaly



## OSR anomaly

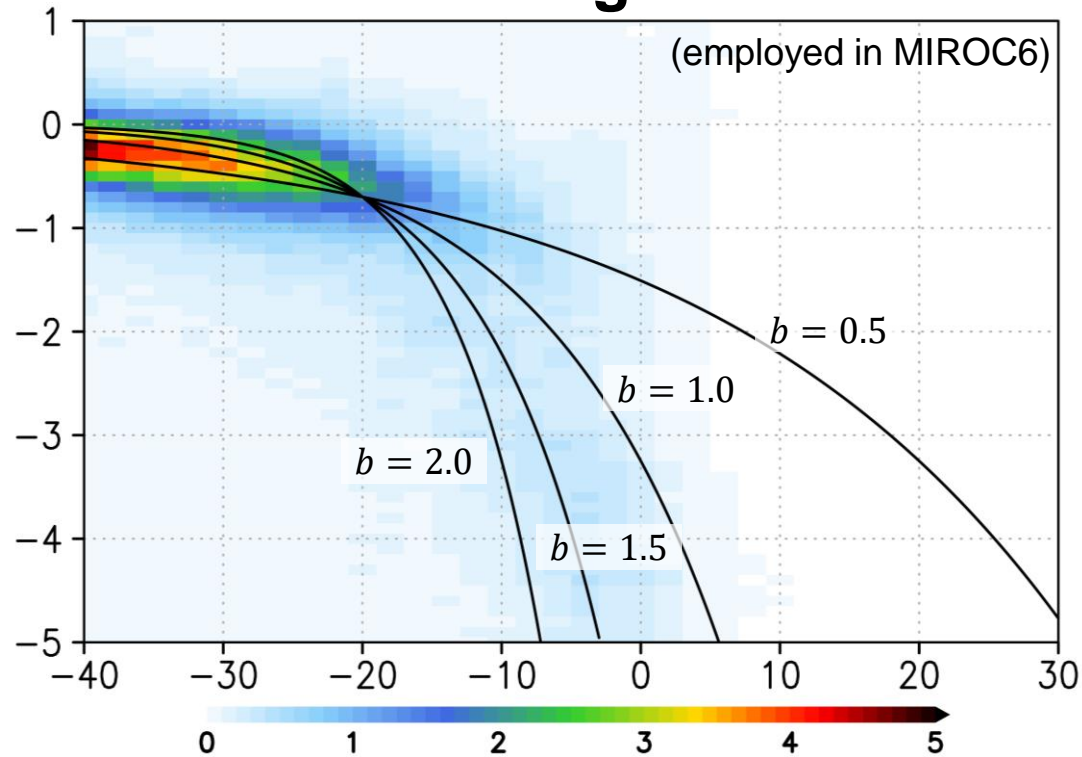


# $Z_e - v_d$ relation on 2dPDF of NICT observation

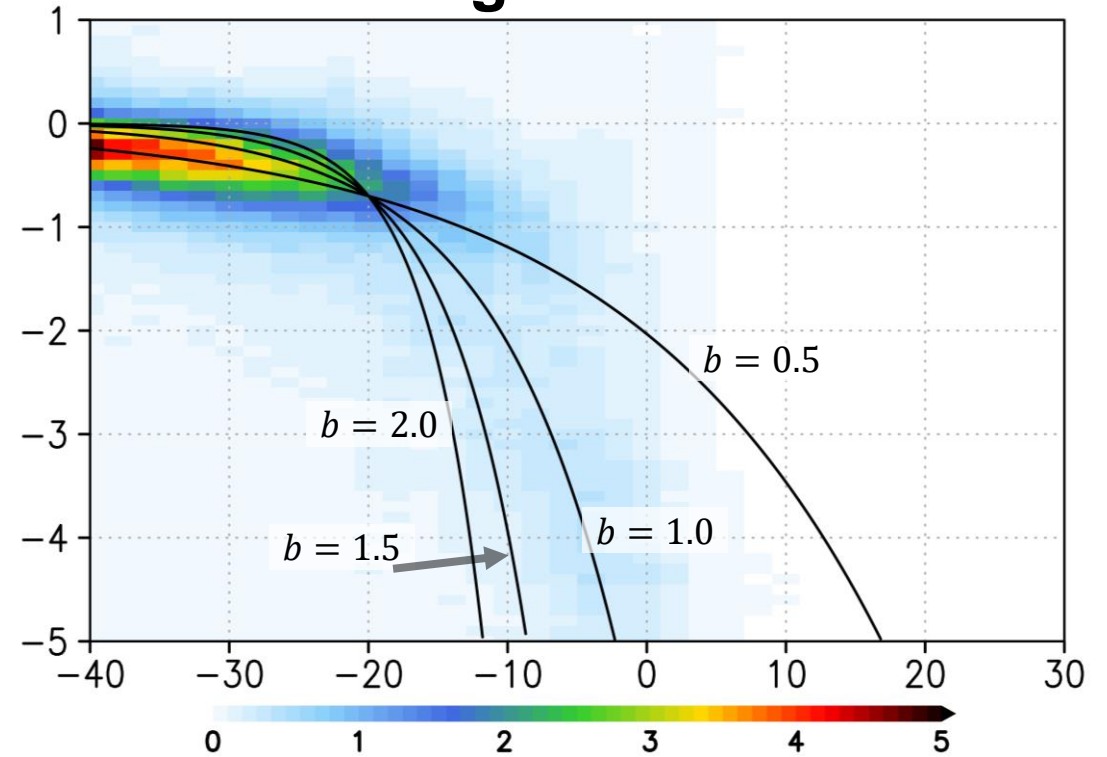
- ✓  $Z_e - v_d$  relation can be made in the similar way as  $Z - R$  relation.
  - with Rayleigh scattering limit and power law for droplet fall velocity

$$v_f = aD^b$$

## modified gamma



## log normal



- This line fitting contributes to estimate the scale  $b$  of droplet fall velocity.
- Different scale is suggested:  $b \approx 0.5$  for snow,  $b \approx 1.0$  or  $1.5$  for rain.

# Summary

## 1. New simulator for doppler velocity is developed on COSP2.

- Droplet fall velocity and vertical air motion are newly handled.

## 2. Distribution of $v_d$ in MIROC6 is not good, actually.

- slower  $v_d$  and smaller its variance, overestimated  $Z_e$
- separated by categories of cloud/precipitation droplet

## 3. Significant impact on climate by tuning of droplet fall velocity

- Large modification of cloud fraction and strong cooling effect.
- Cloud microphysics schemes still have much room for improvement.

## 4. Estimation of scale of fall velocity by $Z_e$ - $v_d$ relation

- Different scale among cloud, rain, snow droplets?
- Implication for shape of ice particles?

- ✓ We are now preparing to publish this new simulator for COSP2.
- ✓ I'd like to see this simulator used with EarthCARE observation to understand model variability and to improve schemes related to clouds.

**Supplement**

# Details of tuning setup

**PL08:  $b_1 - b_2 \exp(-b_3 D) + (b_2 - b_1) \exp(-5b_3 D)$**  or **Power law:  $aD^b$**

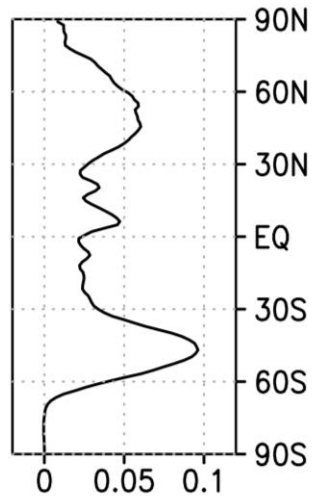
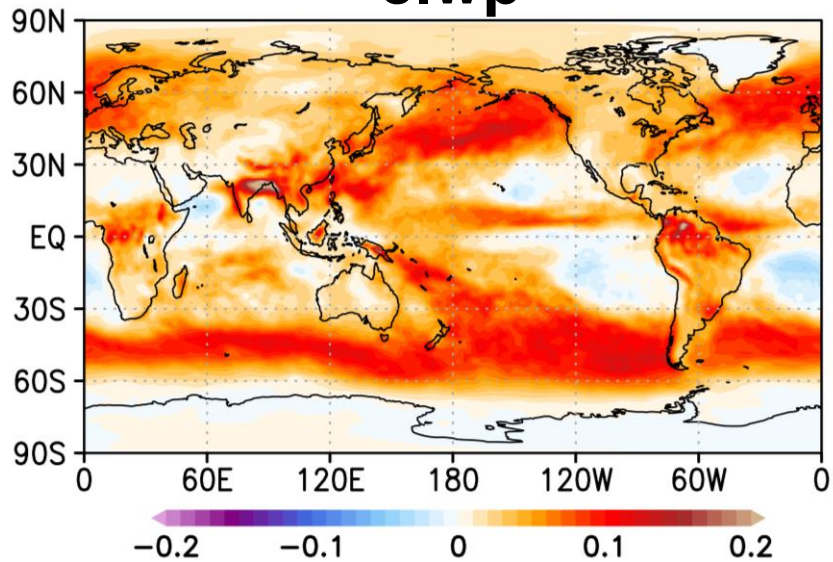
ctrl	large-scale condensation				cumulus convection			
	cloud liq.	cloud ice	rain	snow	cloud liq.	cloud ice	rain	snow
droplet								
viscous drag	×	×	×	×	×	×	○	○
PL08	○	×	○	×	○	×	×	×
Power law; $a$	-	1.107	-	3.321	-	1.107	842	4.84
Power law; $b$	-	0.22	-	0.22	-	0.22	0.8	0.25

**$b_1 = 9.65, b_2 = 10.43, b_3 = 600$**

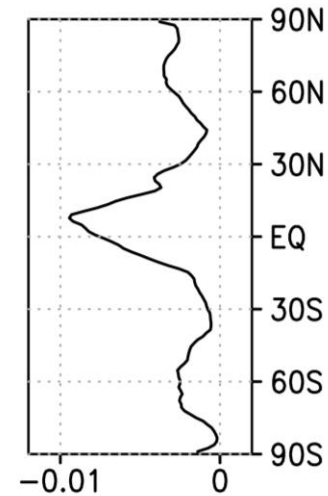
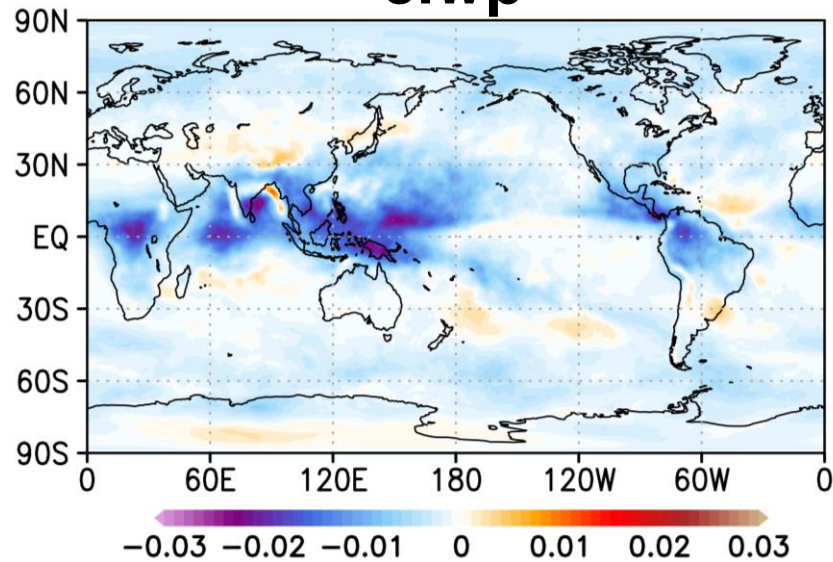
tuning	large-scale condensation				cumulus convection			
	cloud liq.	cloud ice	rain	snow	cloud liq.	cloud ice	rain	snow
droplet								
viscous drag	○	○	○	○	○	○	○	○
PL08	○	×	○	×	○	×	×	×
Power law; $a$	-	300	-	100	-	300	842	4.84
Power law; $b$	-	0.8	-	0.5	-	0.8	0.8	0.25

**$b_1 = 10.0, b_2 = 10.43, b_3 = 1500$**

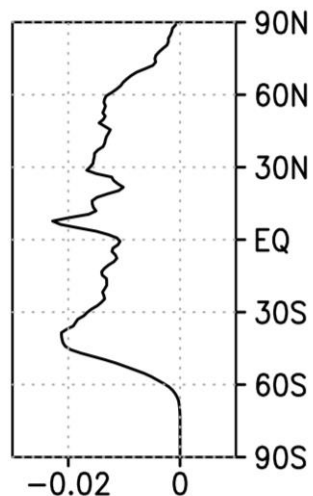
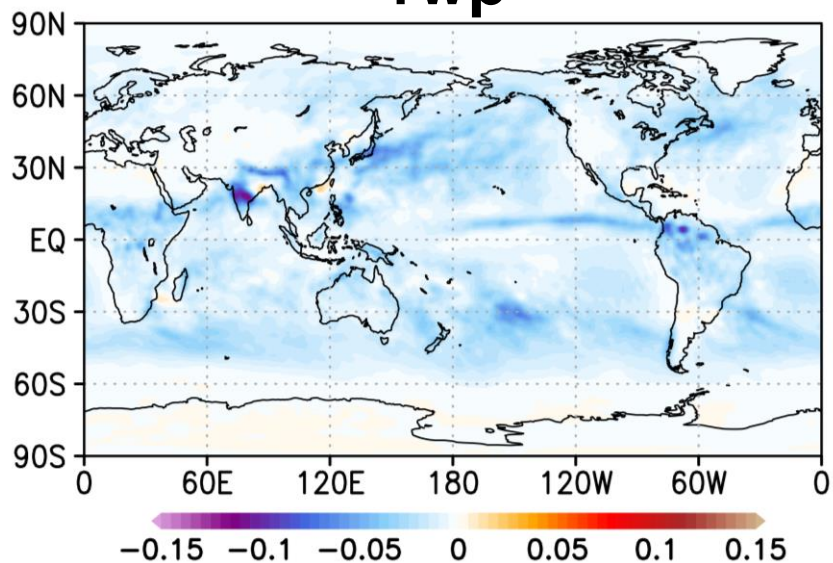
**clwp**



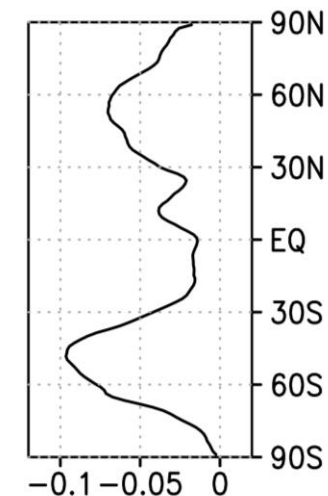
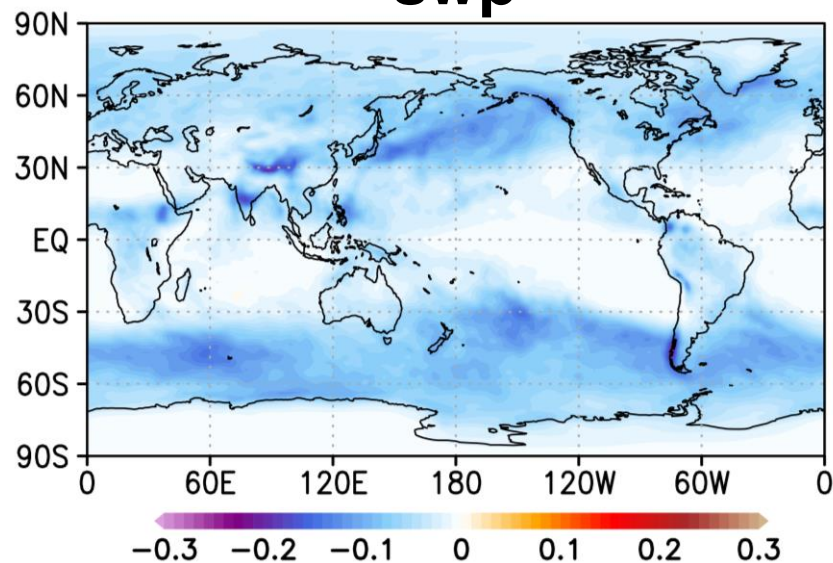
**ciwp**



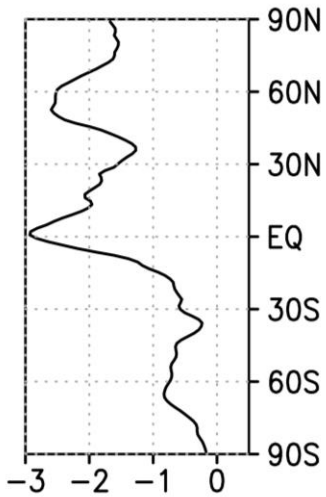
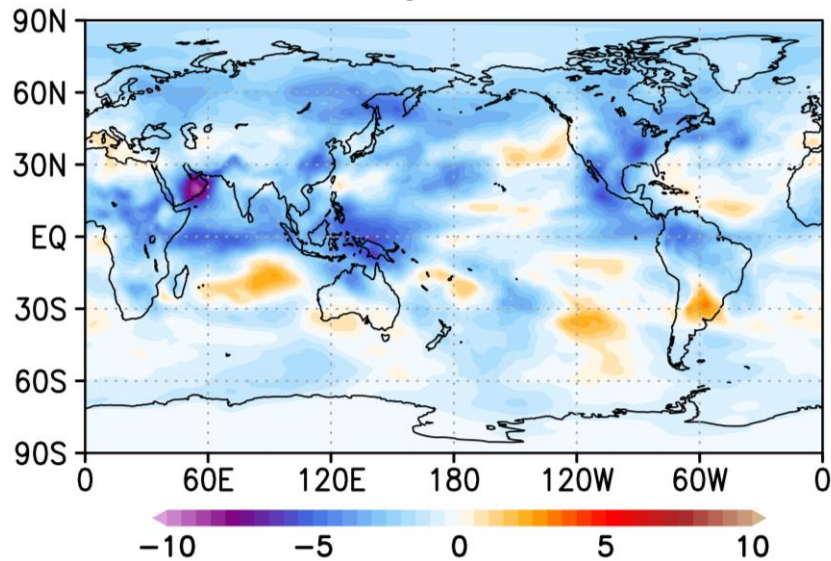
**rwp**



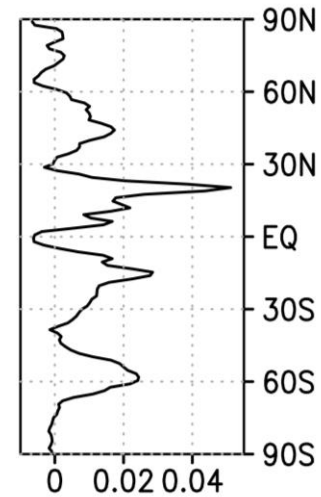
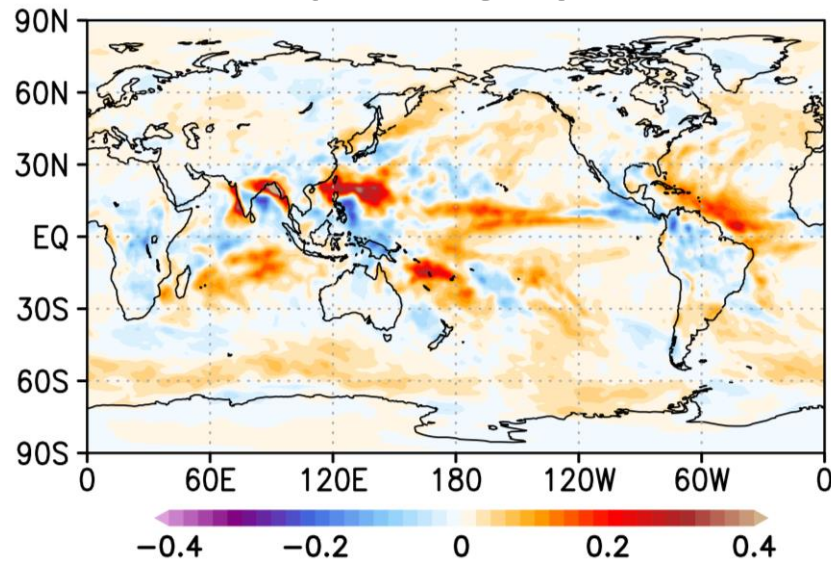
**swp**



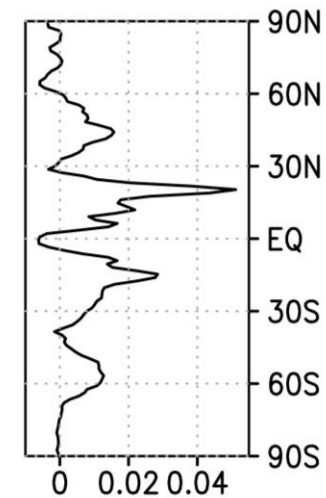
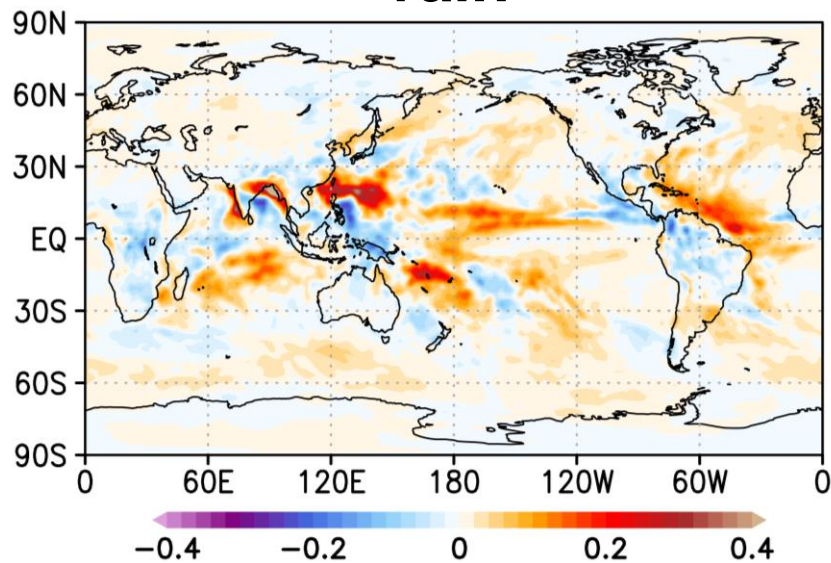
**CWV**



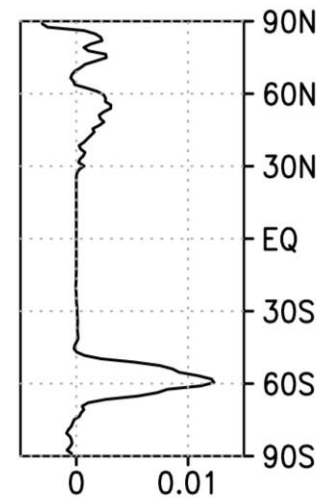
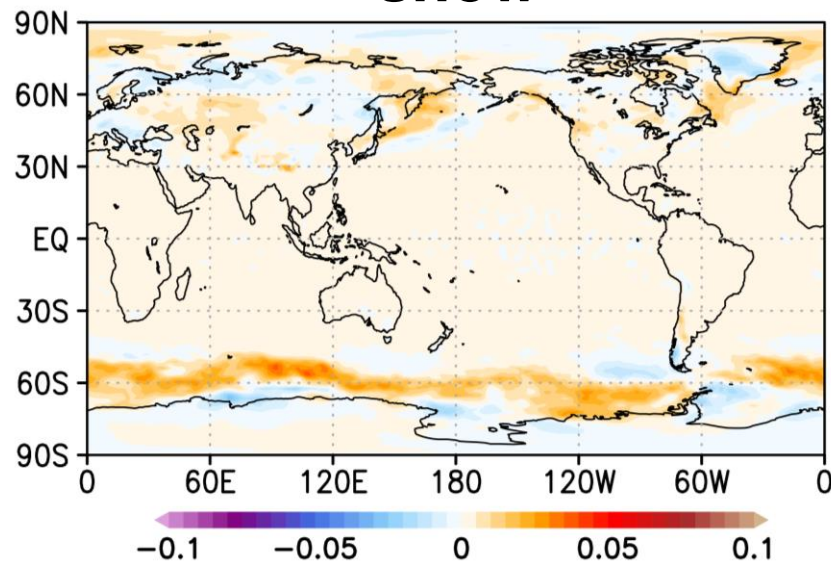
**rain + snow**



**rain**

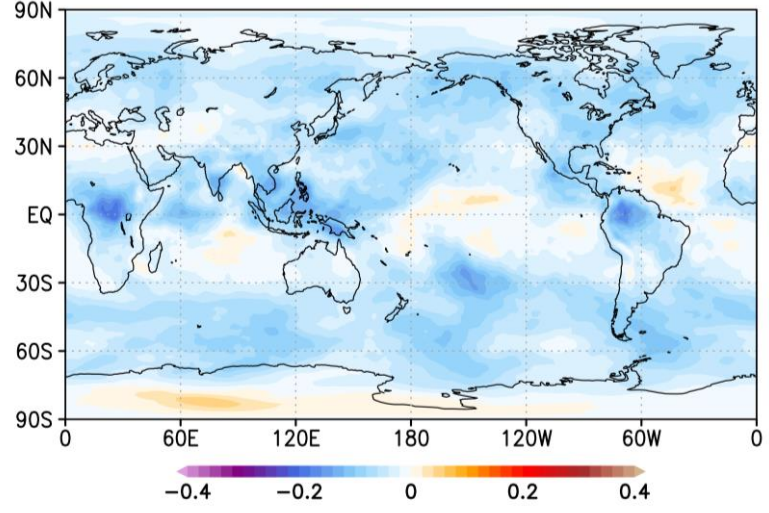


**snow**

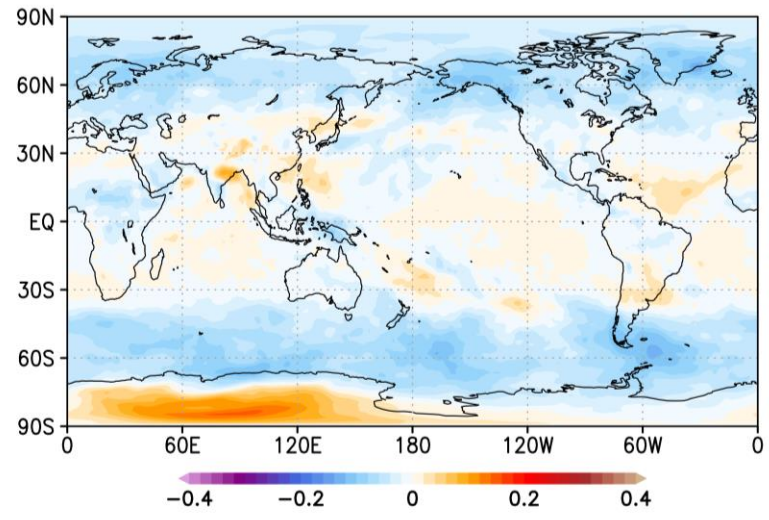




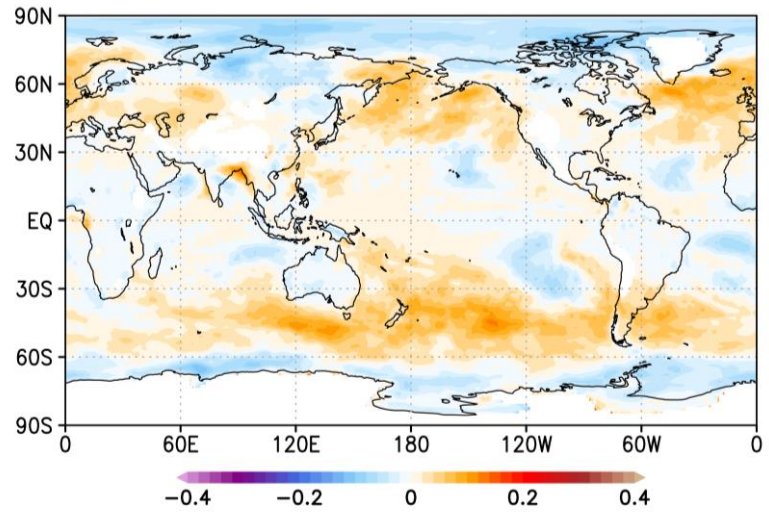
cloud fraction @ 300hPa



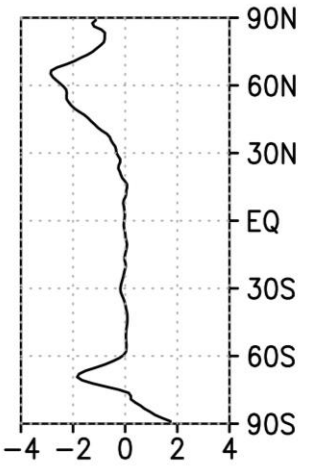
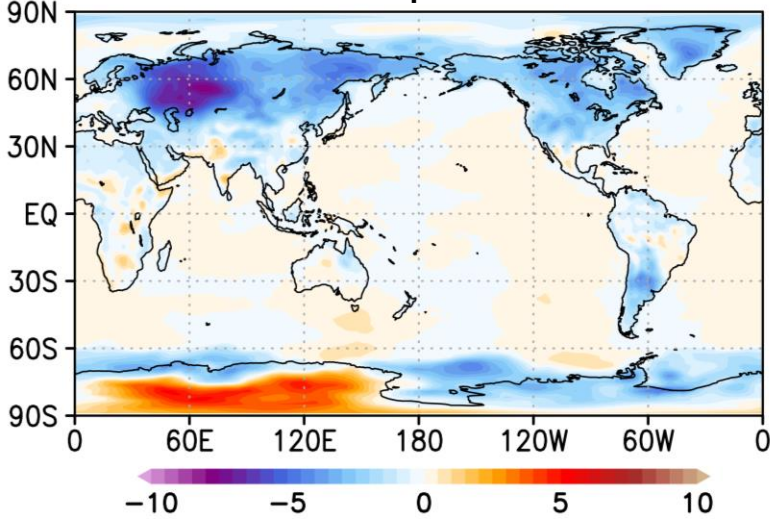
cloud fraction @ 500hPa



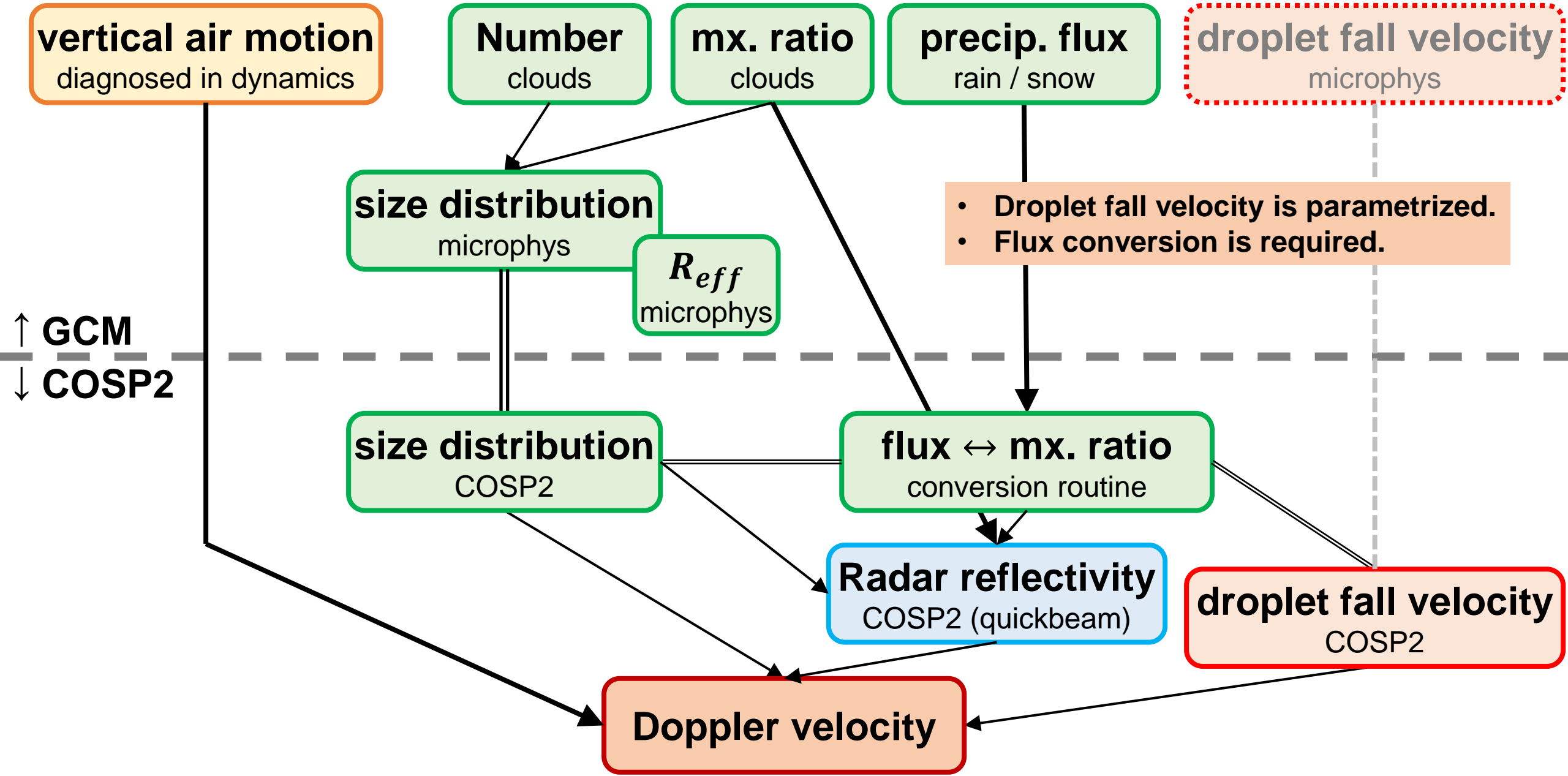
cloud fraction @ 850hPa



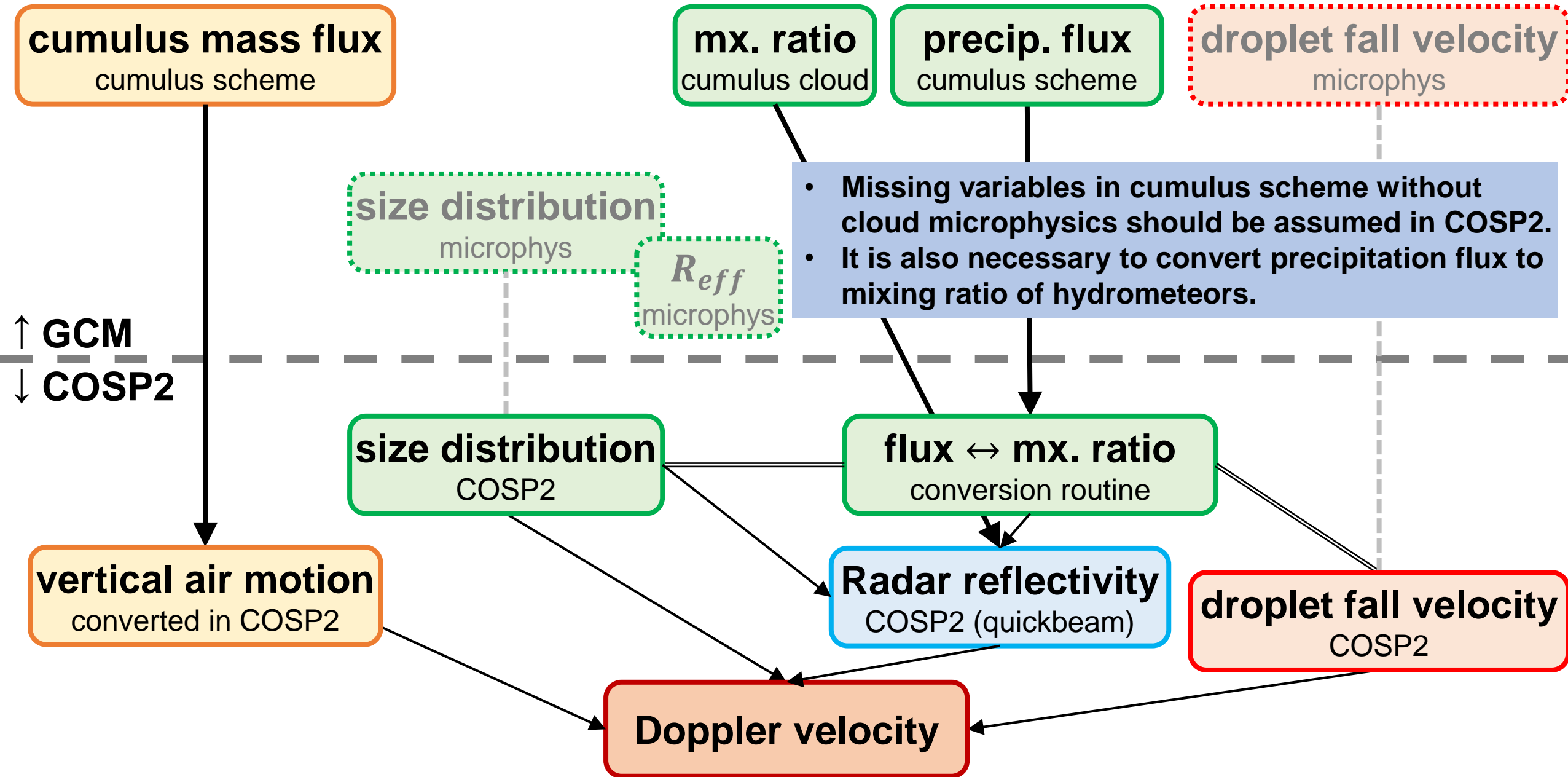
2m temperature



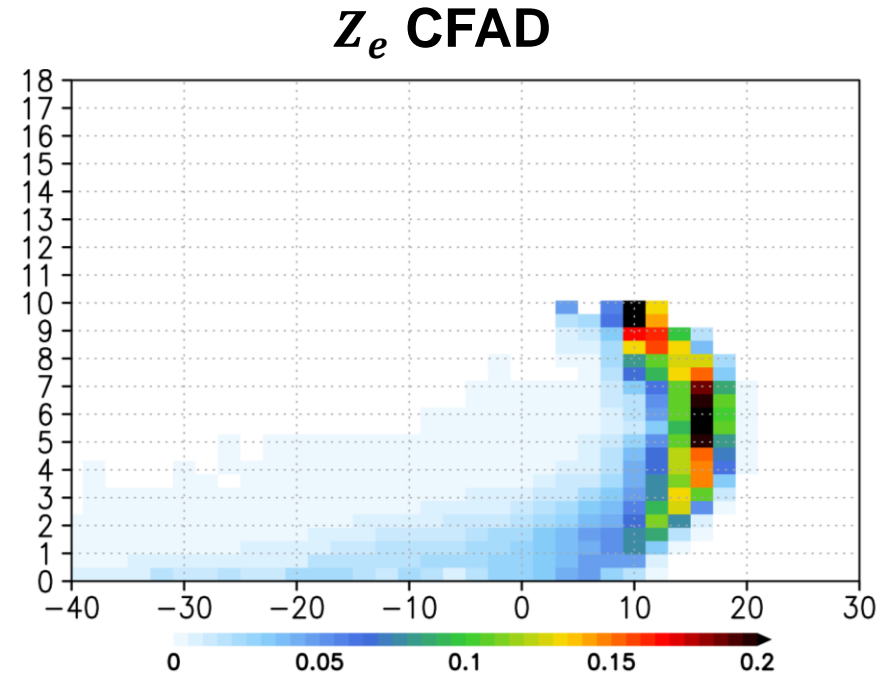
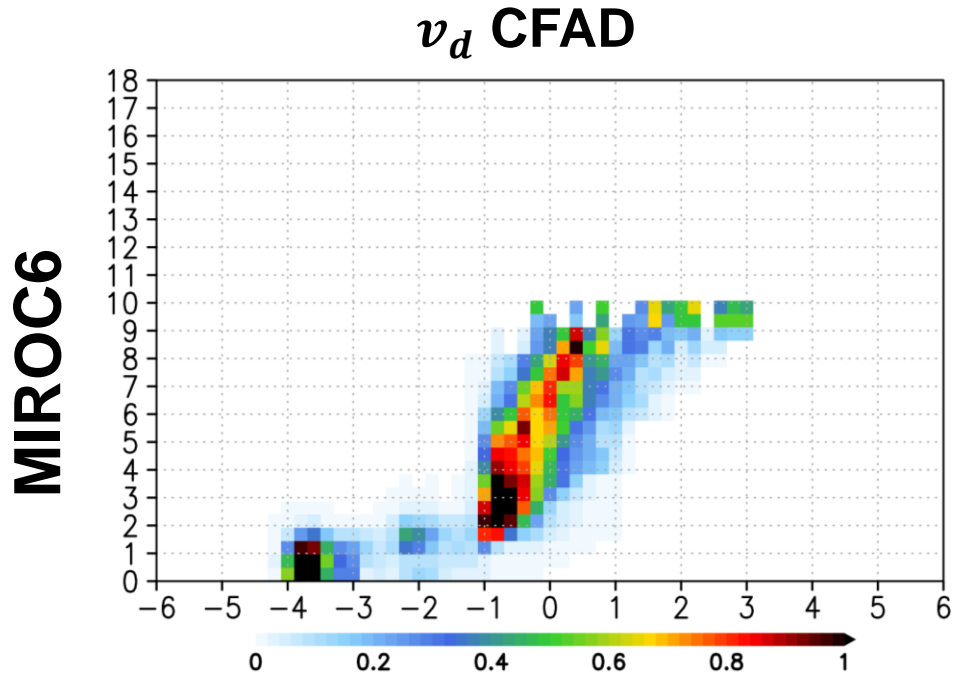
# Simulator design: diagnostic precipitation scheme



# Simulator design: cumulus scheme



# $v_d$ simulator on cumulus precipitation in MIROC6



corresponding area to NICT obs.

$$v_d = \underbrace{w}_{\text{cumulus mass flux}} + \frac{\int_{r_{\min}}^{r_{\max}} \frac{dn(r)}{dr} \cdot C_{bk}(r) v_f(r) dr}{\int_{r_{\min}}^{r_{\max}} \frac{dn(r)}{dr} \cdot C_{bk}(r) dr},$$

Not comparable to NICT monthly statistics.  
We should pick up cumulus cloud case.

$Z_e$  in COSP2 ...

$$Z_e = 10^{18} \frac{\lambda^4}{\pi^5 K} \int C_{bk} n(D) dD$$

$$m = \frac{\pi}{6} \rho N M_3$$

$$\frac{1}{N} = \frac{\pi}{6} \rho \frac{M_3}{m}$$

$$N = \frac{6 m}{\pi \rho M_3}$$

Rayleigh Limit ...

$$C_{bk} = \frac{\pi^5 K}{\lambda^4} D^6$$

$$\rightarrow Z_e = 10^{18} \int D^6 n(D) dD = 10^{18} N M_6$$

$$M_k = \frac{1}{N} \int D^k n(D) dD$$

$$Z_e = 10^{18} \frac{6 M_6 m}{\pi M_3 \rho}$$

$$v_d = \cancel{W} + \frac{\int_{r_{\min}}^{r_{\max}} dn(r)/dr \cdot C_{bk}(r) v_f(r) dr}{\int_{r_{\min}}^{r_{\max}} dn(r)/dr \cdot C_{bk}(r) dr},$$

$$v_f = a D^b$$

$$v_d = \frac{a M_{6+b}}{M_6}$$

$$dBZ_e = 10 \log_{10} Z_e$$

$$Z_e = 10^{\frac{dBZ_e}{10}}$$

mod. gamma

$$n(D) = \frac{N}{\Gamma(\mu)D} \left(\frac{D}{D_0}\right)^\mu \exp\left(-\frac{D}{D_0}\right)$$

$$NM_k = \int D^k n(D) dD$$

$$M_k = D_0^k \frac{\Gamma(\mu + k)}{\Gamma(\mu)}$$

$$\begin{aligned} v_d &= \frac{aM_{6+b}}{M_6} \\ &= aD_0^b \frac{\Gamma(\mu + 6 + b)}{\Gamma(\mu + 6)} \end{aligned}$$

$$D_0^b = \frac{v_d}{a} \frac{\Gamma(\mu + 6)}{\Gamma(\mu + 6 + b)}$$

$$\begin{aligned} v_d &= aM_{6+b} 10^{18} \frac{1}{Z_e} \frac{6}{\pi} \frac{m}{M_3} \\ &= \frac{6 \times 10^{18}}{\pi} am \frac{M_{6+b}}{M_3} \frac{1}{Z_e} \\ &= \frac{6 \times 10^{18}}{\pi} aD_0^{3+b} m \frac{\Gamma(\mu + 6 + b)}{\Gamma(\mu + 3)} \frac{1}{Z_e} \end{aligned}$$

$$\begin{aligned} Z_e &= 10^{18} \frac{6}{\pi} \frac{M_6}{M_3} \frac{m}{\rho} \\ &= 10^{18} \frac{6}{\pi} \frac{m}{\rho} D_0^3 \frac{\Gamma(\mu + 6)}{\Gamma(\mu + 3)} \end{aligned}$$

$$D_0^3 = 10^{-18} \frac{\pi}{6} \frac{\rho}{m} \frac{\Gamma(\mu + 3)}{\Gamma(\mu + 6)} Z_e$$

$$v_d = a \frac{\Gamma(\mu + 6 + b)}{\Gamma(\mu + 6)} \left( 10^{-18} \frac{\pi}{6} \frac{\rho}{m} \frac{\Gamma(\mu + 3)}{\Gamma(\mu + 6)} Z_e \right)^{\frac{b}{3}}$$

log normal

$$n(D) = \frac{1}{D} \frac{N}{\sqrt{2\pi} \log \sigma} \exp \left[ -\frac{(\log D - \log \mu)^2}{2(\log \sigma)^2} \right]$$

$$NM_k = \int D^k n(D) dD$$

$$M_k = \exp \left( k \log \mu + \frac{k^2 (\log \sigma)^2}{2} \right)$$

$$2R_{eff} = \frac{M_3}{M_2} = \mu \exp \left[ \frac{5}{2} (\log \sigma)^2 \right]$$

$$(\log \sigma)^2 = \frac{2}{5} \log \left( \frac{2R_{eff}}{\mu} \right)$$

$$\begin{aligned} M_k &= \exp \left( k \log \mu + \frac{k^2}{5} \log \left( \frac{2R_{eff}}{\mu} \right) \right) \\ &= \mu^k \cdot \left( \frac{2R_{eff}}{\mu} \right)^{\frac{k^2}{5}} = \mu^{k - \frac{k^2}{5}} D_{eff}^{\frac{k^2}{5}} \end{aligned}$$

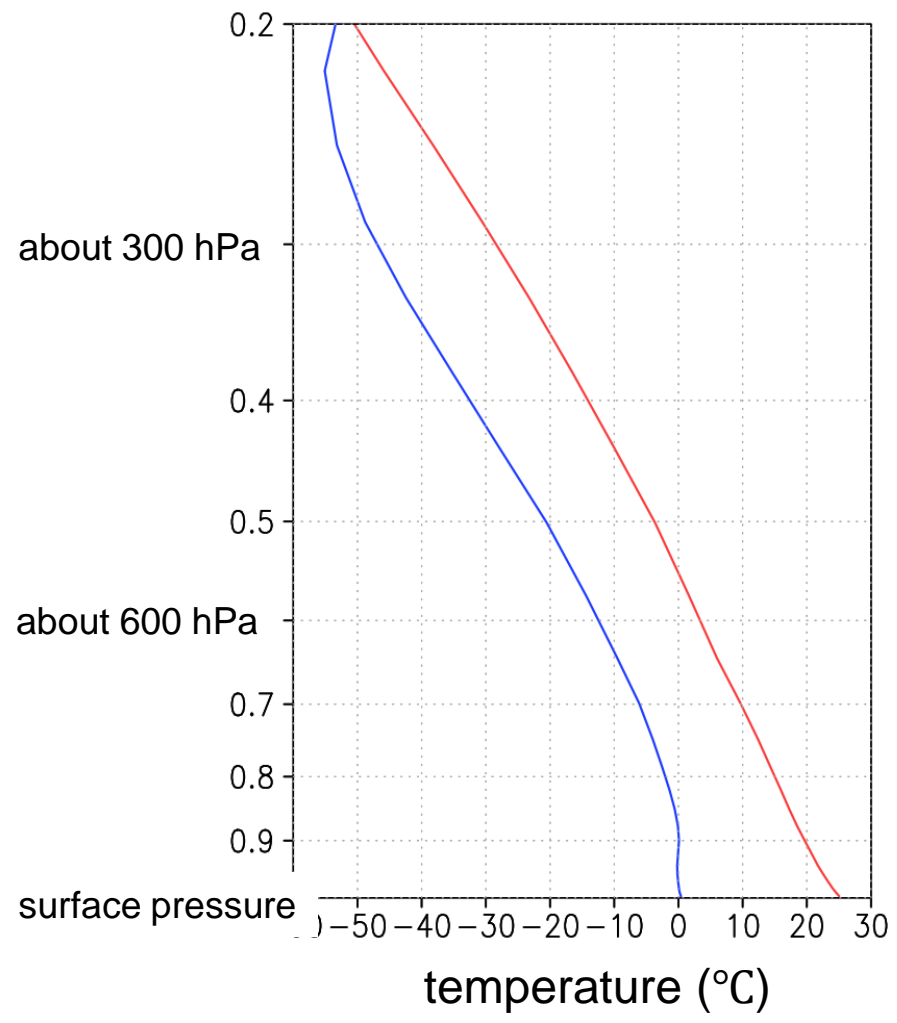
$$\begin{aligned} v_d &= \frac{aM_{6+b}}{M_6} = a \mu^{b - \frac{(6+b)^2 - 6^2}{5}} (D_{eff})^{\frac{(6+b)^2 - 6^2}{5}} \\ &= a \mu^{-\frac{b(b+7)}{5}} D_{eff}^{\frac{b(b+12)}{5}} \end{aligned}$$

$$Z_e = 10^{18} \frac{6 M_6 m}{\pi M_3 \rho} = 10^{18} \frac{6 m}{\pi \rho} \mu^{-\frac{12}{5}} (D_{eff})^{\frac{27}{5}}$$

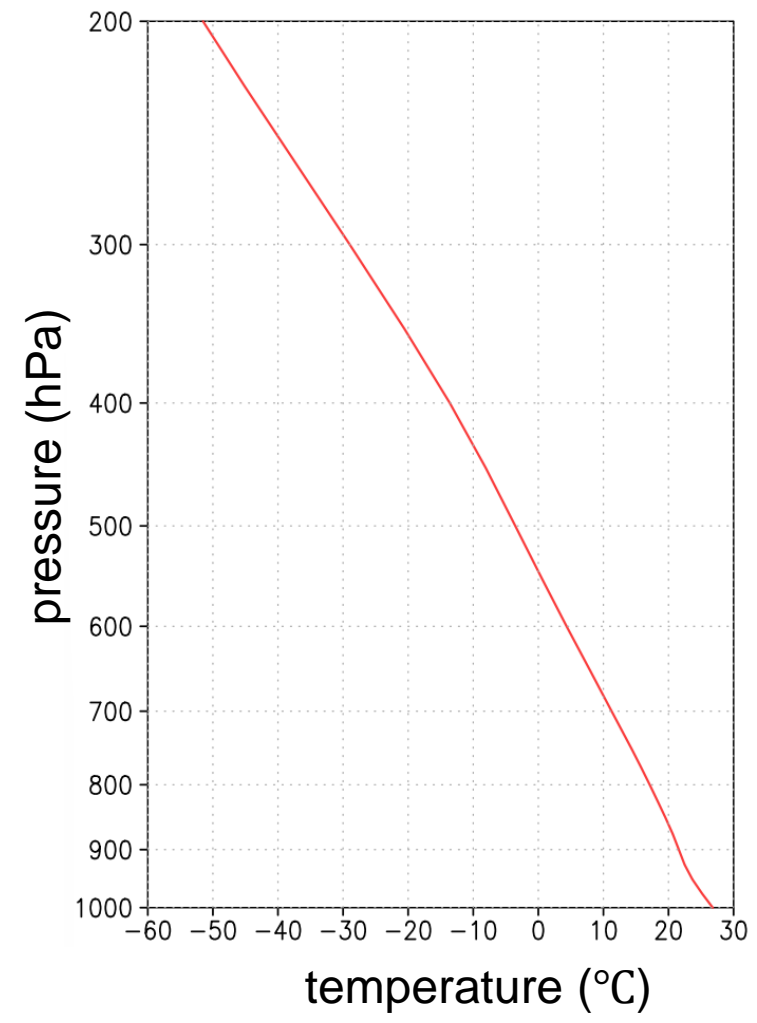
$$\begin{aligned} v_d &= a \mu^{-\frac{b(b+7)}{5}} \left( 10^{-18} \frac{\pi \rho}{6 m} \mu^{\frac{12}{5}} Z_e \right)^{\frac{b(b+12)}{27}} \\ &= a \mu^{-\frac{b(b+3)}{9}} \left( 10^{-18} \frac{\pi \rho}{6 m} \right)^{\frac{b(b+12)}{27}} Z_e^{\frac{b(b+12)}{27}} \end{aligned}$$

$$\begin{aligned} v_d &= a D_{eff}^{\frac{b(b+12)}{5}} \left( 10^{-18} \frac{\pi \rho}{6 m} D_{eff}^{-\frac{27}{5}} Z_e \right)^{\frac{b(b+7)}{12}} \\ &= a D_{eff}^{-\frac{b(b+3)}{4}} \left( 10^{-18} \frac{\pi \rho}{6 m} \right)^{\frac{b(b+7)}{12}} Z_e^{\frac{b(b+7)}{12}} \end{aligned}$$

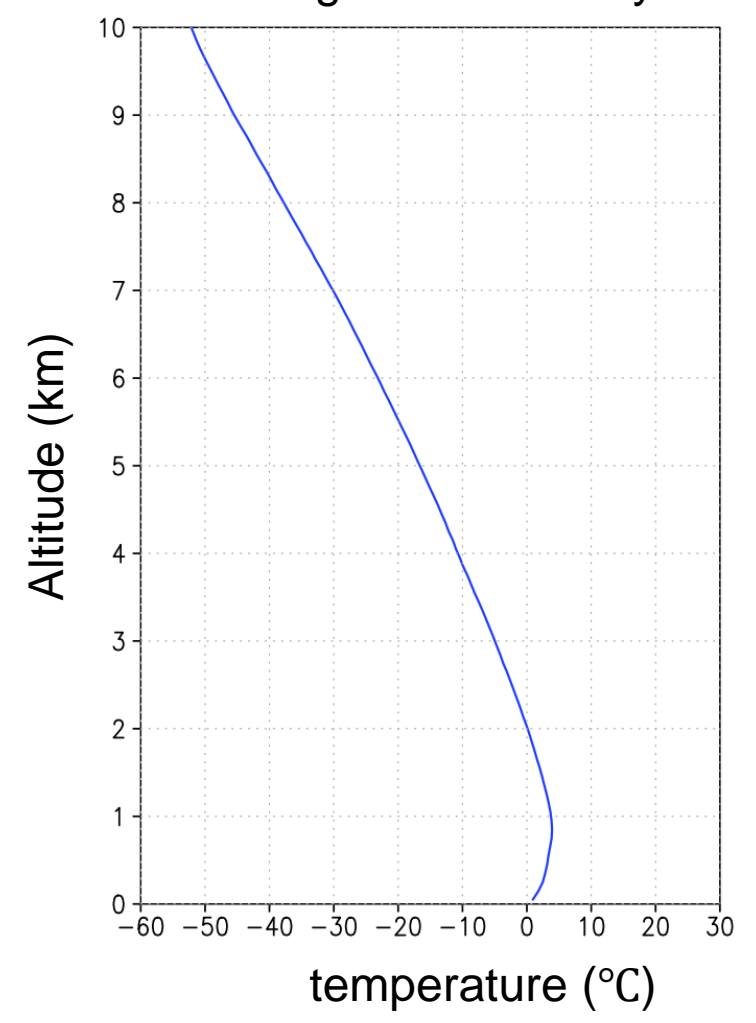
MIROC6 ctrl  
red: NICT  
blue: MOSAiC



JRA3Q (reanalysis)  
NICT  
Aug. 2020 monthly



MOSAic  
Sounding  
Aug. 2020 monthly







# Simulator design: \*\*\*\*\*

