

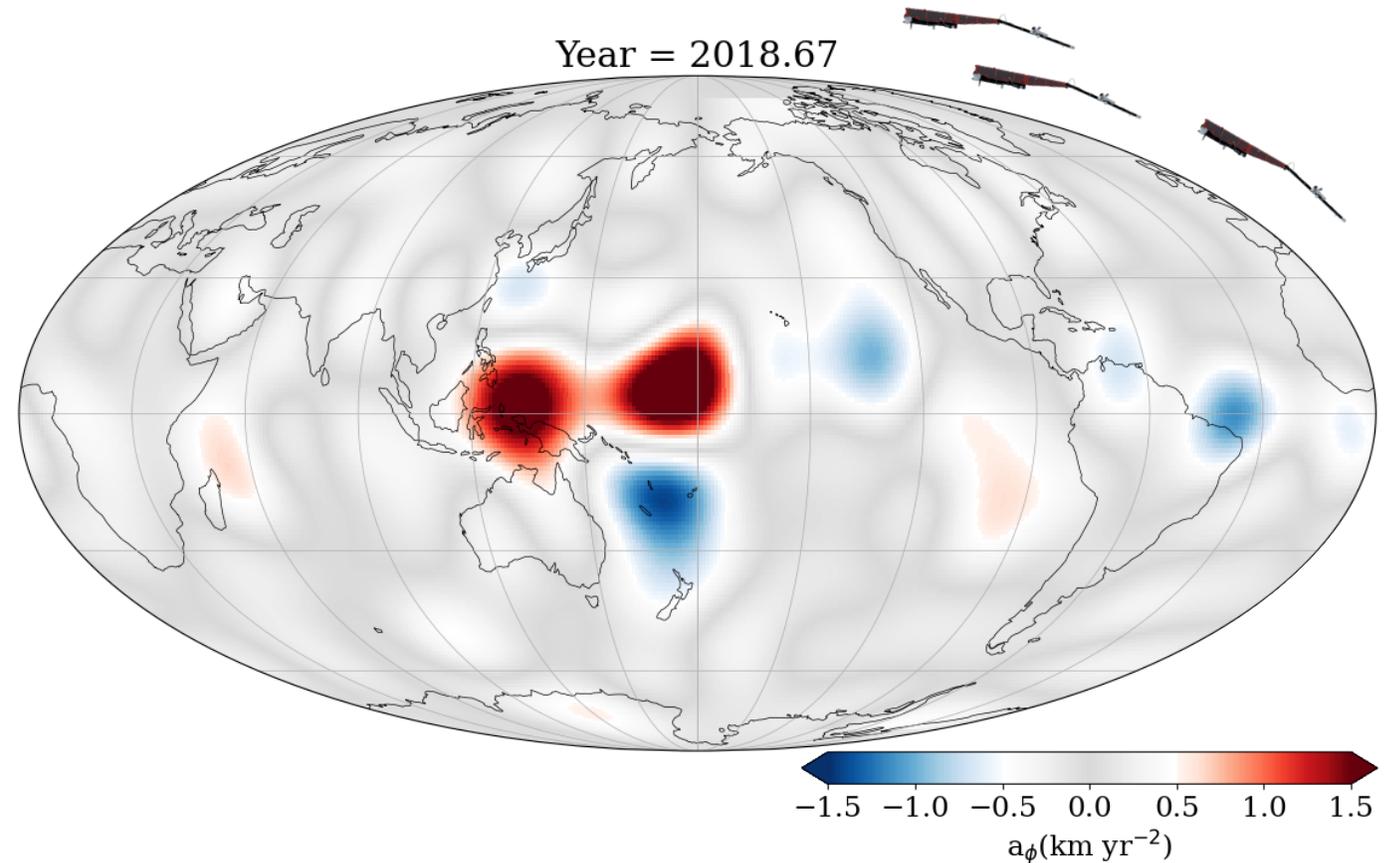


Swarm 10 Year Anniversary
& Science Conference 2024

A Swarm of jerks

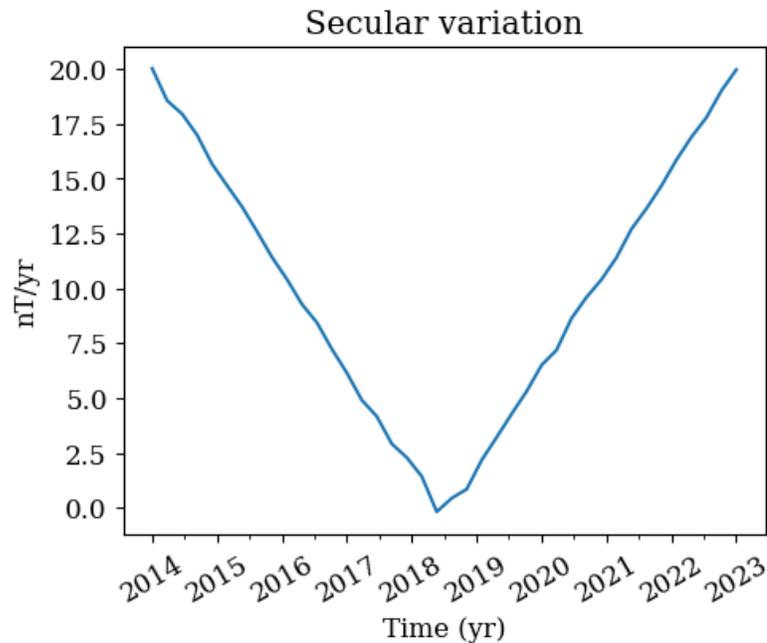
Using the spatial gradient tensor
for core-surface flow modelling

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Ciarán Beggan², Will Brown², and Richard Holme³

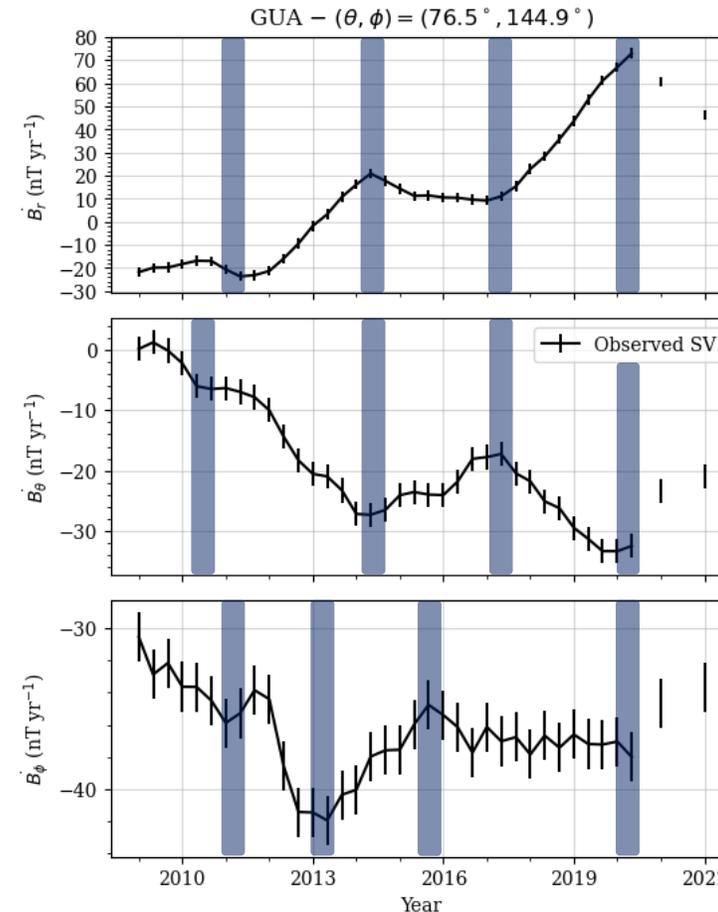


What are geomagnetic jerks?

- The internal magnetic field from the outer core is highly dynamic.
- Geomagnetic jerks: **abrupt*** changes in **secular variation (SV)**, often V or Λ shaped.



Synthetic geomagnetic jerk



Geomagnetic jerks at GUA

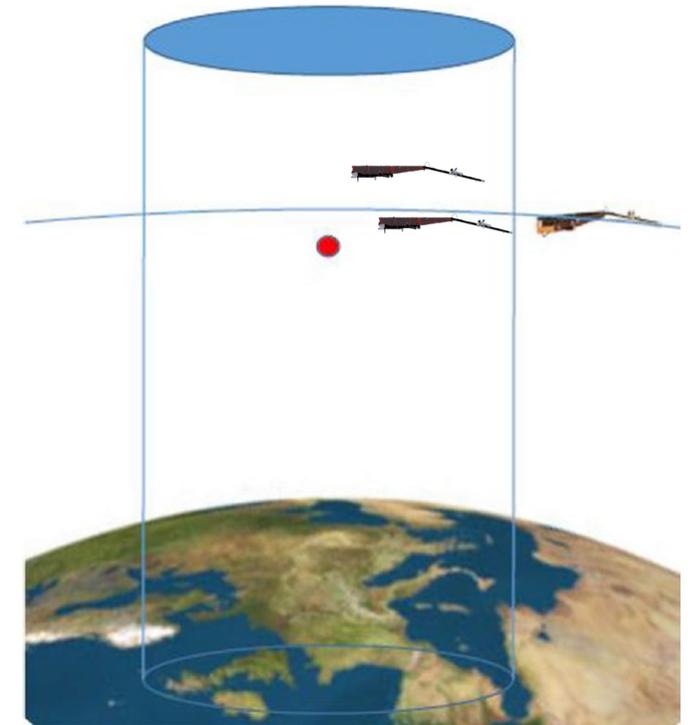
* < 1 year

Geomagnetic virtual observatory (GVO) gradients

- Bin satellite data onto a semi-regular grid on a sphere,
 - Each grid point is the centre of a cylinder (radius = 700 km)
- This is the **GVO**
- By taking along- and across* track differences of the data, we can estimate the **spatial gradients tensor** (Hammer et al. 2022)

$$\begin{pmatrix} [\mathbf{B}]_{rr} & & \\ [\mathbf{B}]_{r\theta} & [\mathbf{B}]_{\theta\theta} & \\ [\mathbf{B}]_{r\phi} & [\mathbf{B}]_{\theta\phi} & [\mathbf{B}]_{\phi\phi} \end{pmatrix}$$

- We can obtain gradients from **Swarm** and **CHAMP**

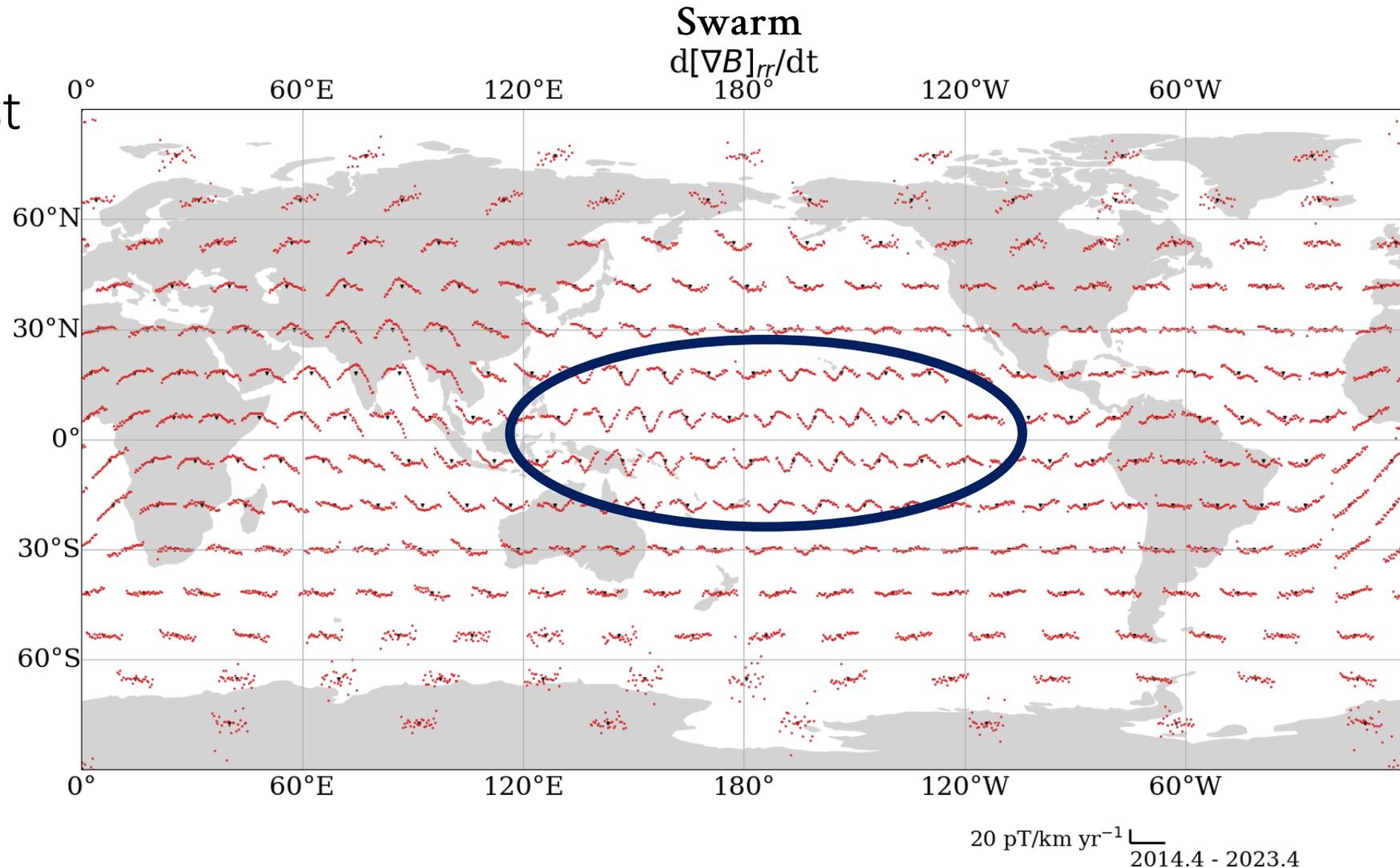


Hammer et al. (2021), EPS

*Across track differences only available with Swarm

GVO gradients from Swarm

- We obtain **SV** from first annual differences
- We use **4-monthly** mean gradient tensor SV
- For **Swarm**, this gives 28 epochs, from **2014.4—2023.4**
- For **CHAMP**, this gives 27 epoch, from **2001.4—2010.0**



Thanks to Jonas Bregnhøj Lauridsen and Chris Finlay at DTU for these data.

Flow inversion – inverting the induction equation

- Assume that magnetic field is “frozen” into the fluid flow → Neglect diffusion
→i.e. “When the field moves, it is moved by flow in the outer core”

- This leads to reduced induction equation:

$$\begin{array}{c} \text{Horizontal surface flow} \\ \dot{B}_r + \nabla_{\mathbf{H}} \cdot (\mathbf{u}_{\mathbf{H}} B_r) = 0 \\ \text{Radial secular variation} \quad \text{Radial magnetic field} \end{array} \equiv$$

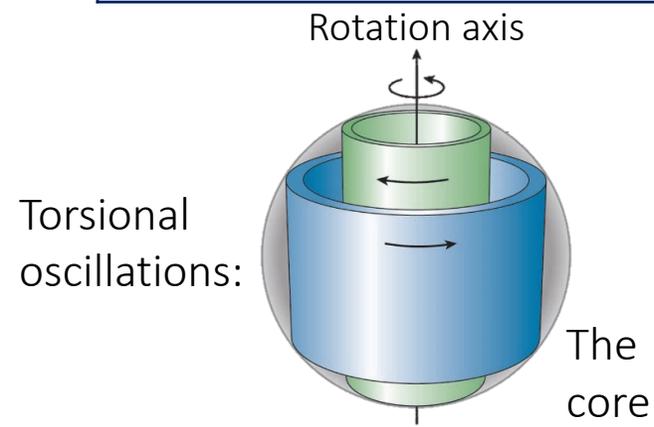
Secular variation is driven by the interaction between flow and the magnetic field

- By decomposing the flow into poloidal (\mathcal{S}) and toroidal (\mathcal{T}) components, we expand data and flow components into spherical harmonics (SH)
- We obtain our model with a damped least squares inversion.
 - Our models are damped: to minimise acceleration between epochs, and to minimise spatial complexity.

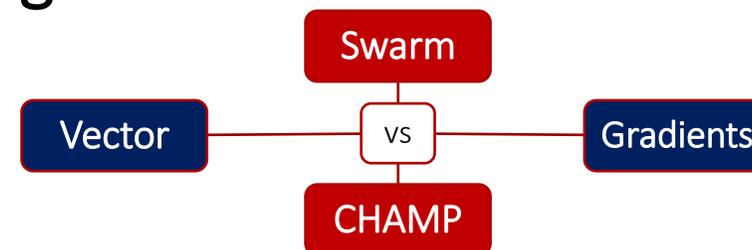
Flow inversion – minimum acceleration and TO-like

- We **obtain our model** with a **damped** least squares inversion.
 - Our models are damped: to **minimise acceleration** between epochs, and to **minimise spatial complexity**.
 - **No use of numerical or stochastic models**, and **without enforcing any flow-geometry** (such as quasi-geostrophy or equatorial symmetry)
- We create two types of models:

Model	Damping
Minimal acceleration	Damp all flow coefficients
Torsional oscillation (TO)-like	No temporal damping on equatorially symmetric zonal flow coefficients

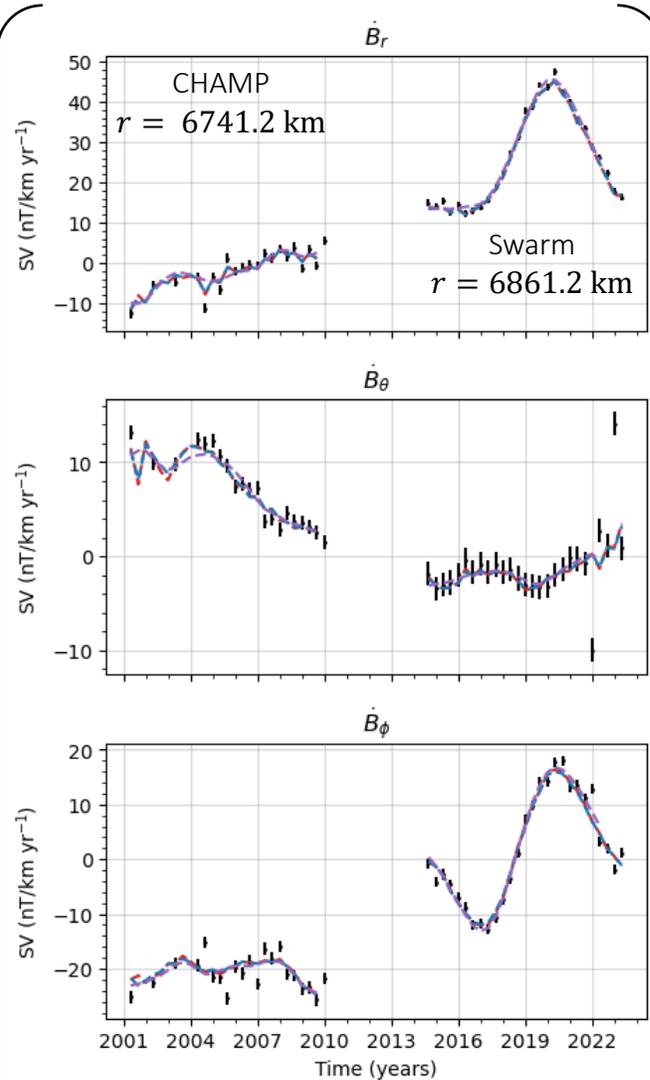


We **create** (and **compare!**) flow models from **vector** and **spatial gradient** data from **CHAMP** and **Swarm**

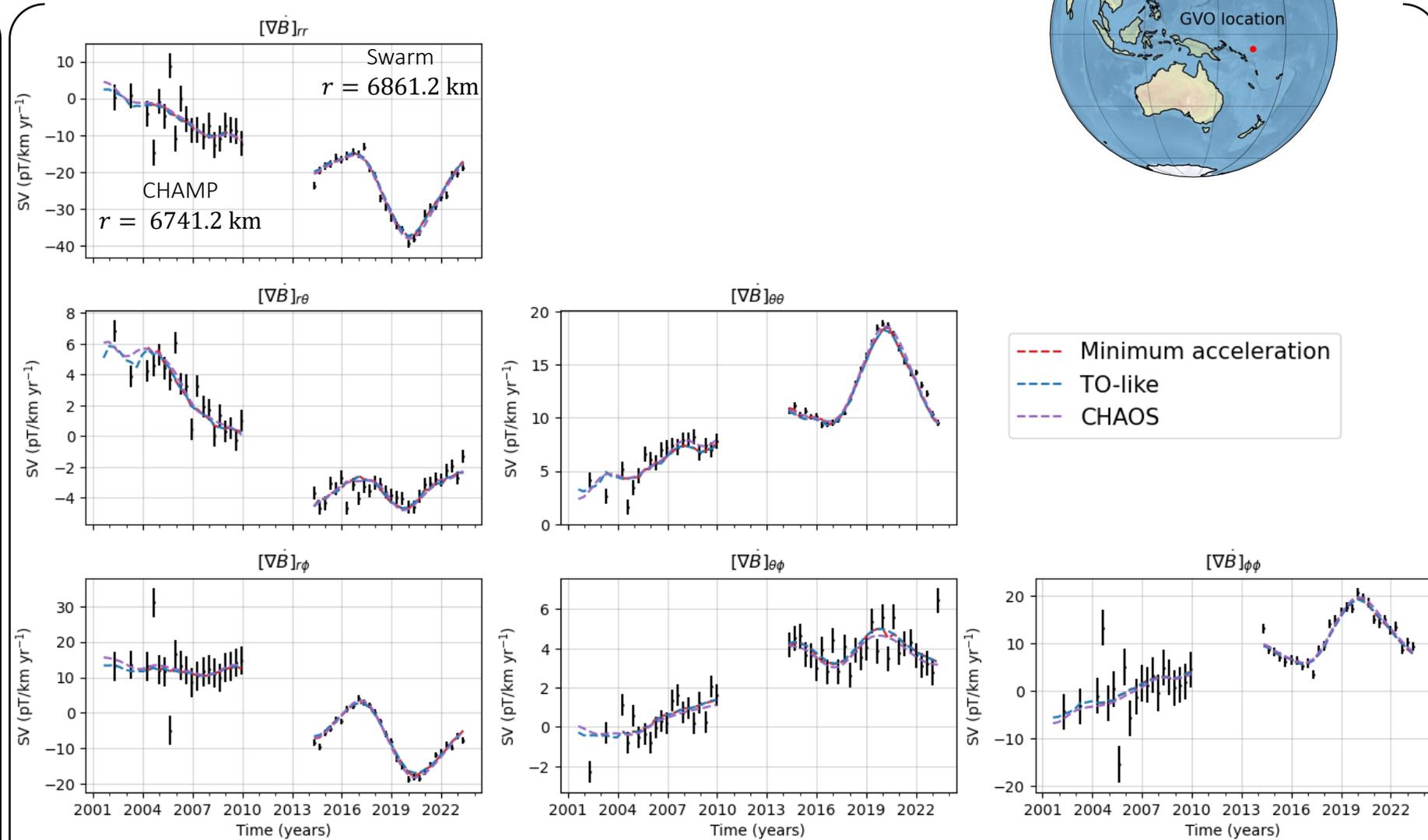


Model performance – good agreement with data

Vector data
RMS misfit = 1.30



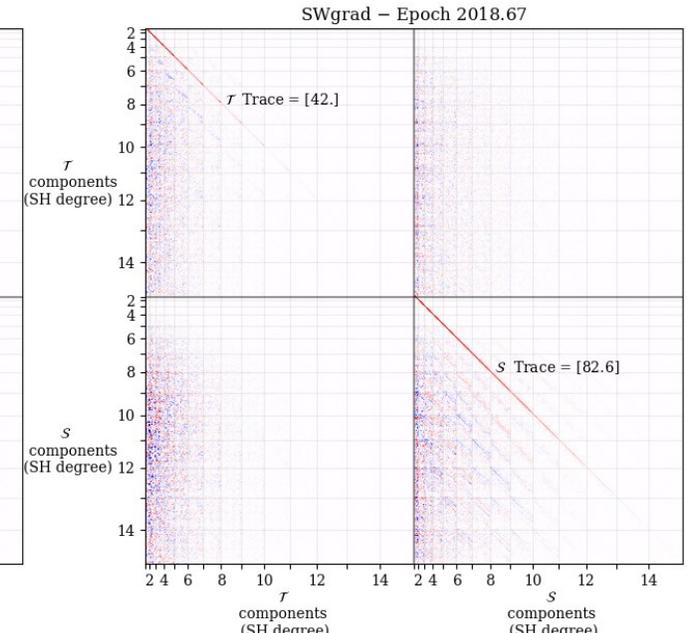
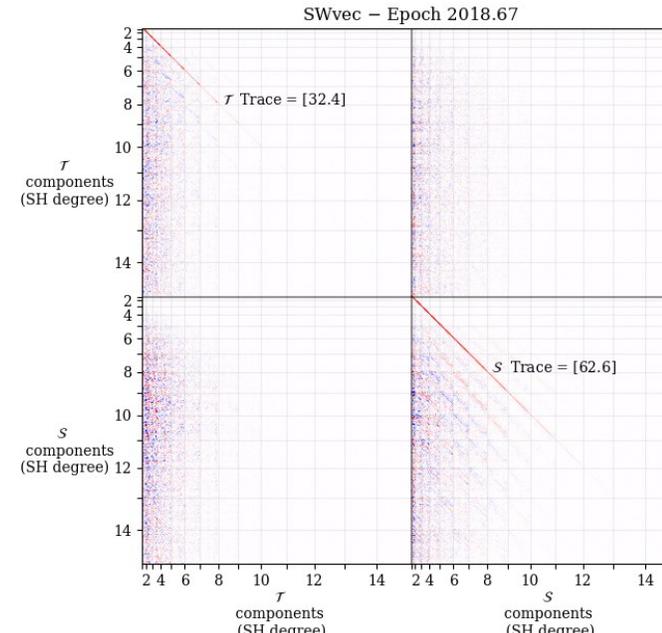
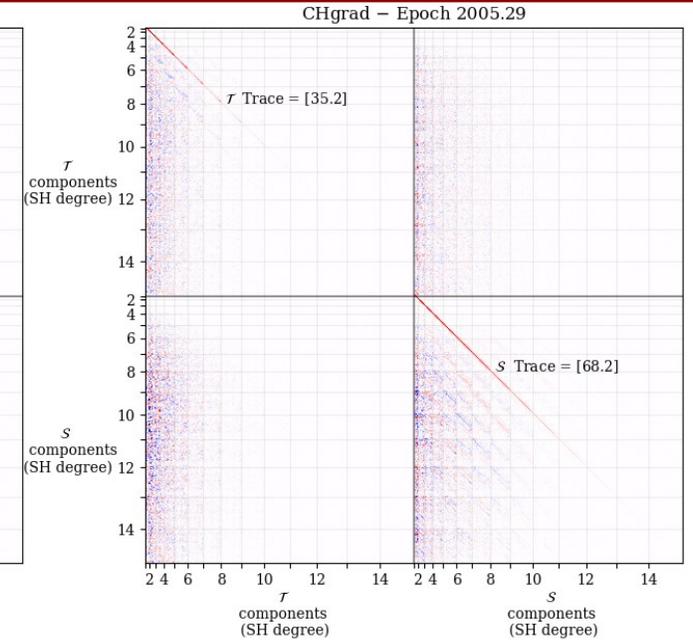
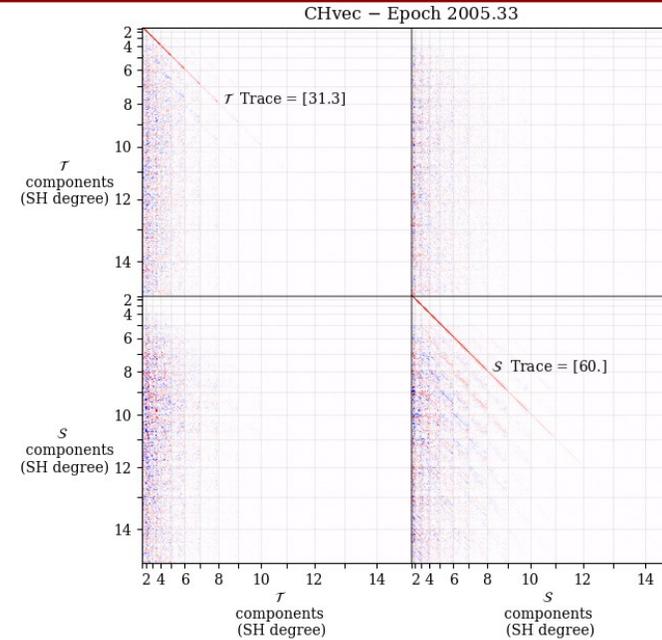
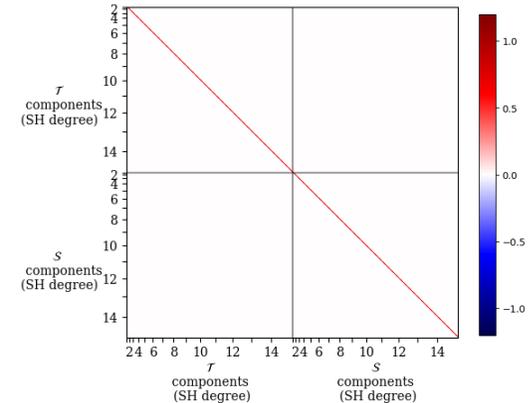
Gradient data
RMS misfit = 1.00



Model resolution – gradients resolve more flow coefficients

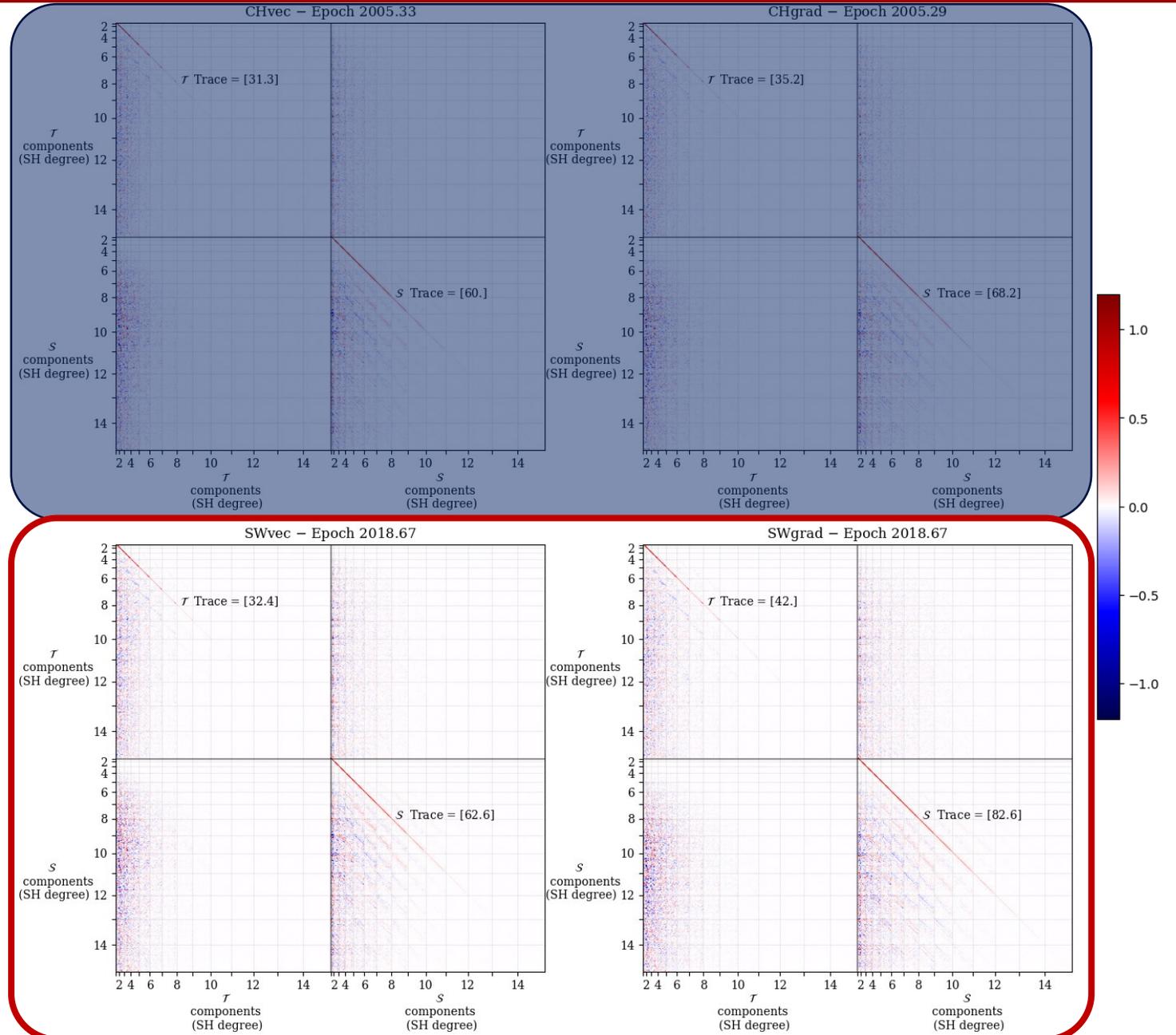
- Resolution matrix shows how well the velocity coefficients are resolved by data
- Trace = sum of diagonals
- Ideal trace for flow of maximum SH degree 14 = 224

Perfect resolution matrix

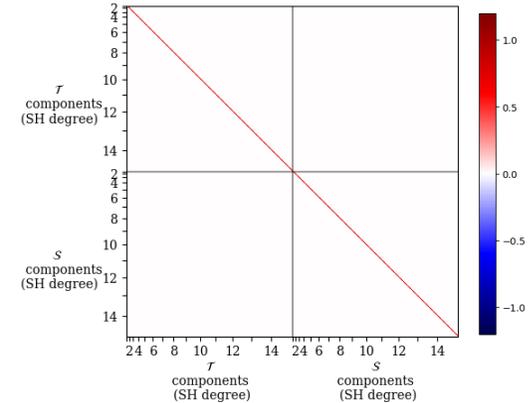


Model resolution – gradients resolve more flow coefficients

- Gradient based models resolve ~133% of that of the vector derived models for Swarm

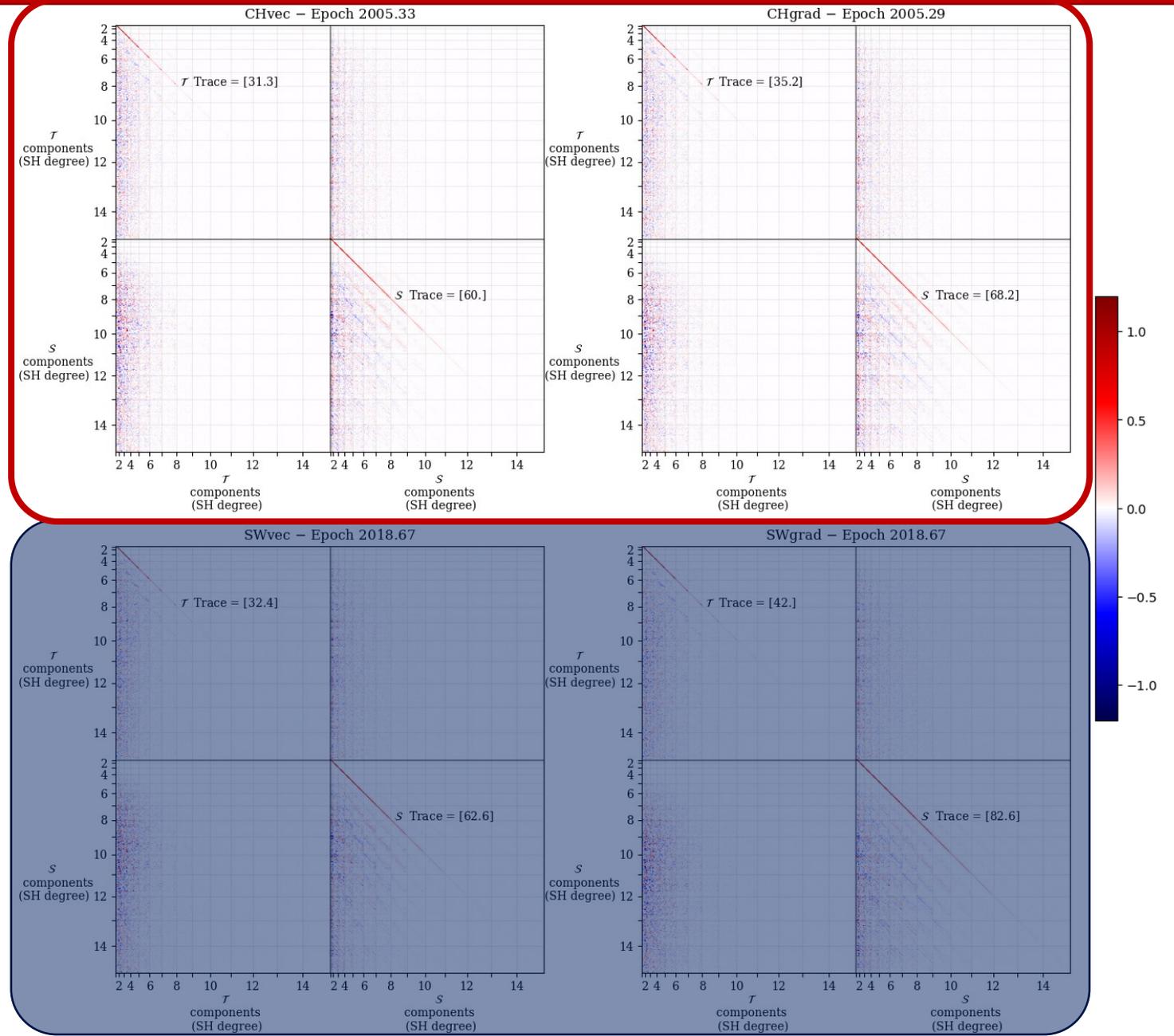


Perfect resolution matrix

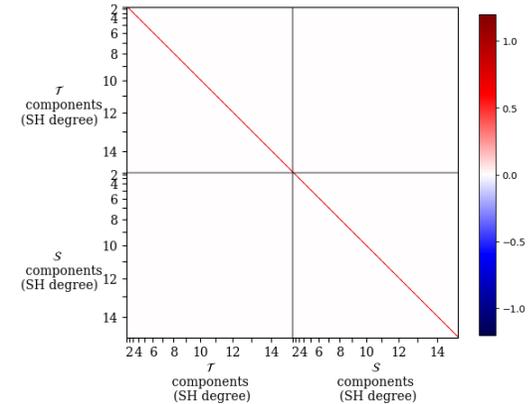


Model resolution – gradients resolve more flow coefficients

- Slight improvement when using gradient data for flow-modelling in CHAMP era

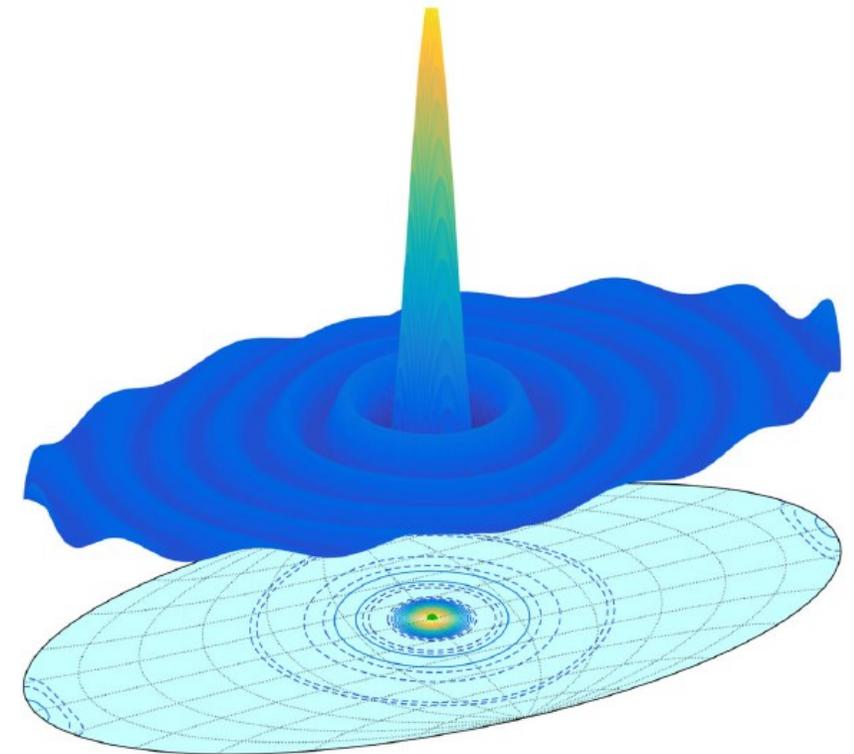


Perfect resolution matrix



Averaging functions

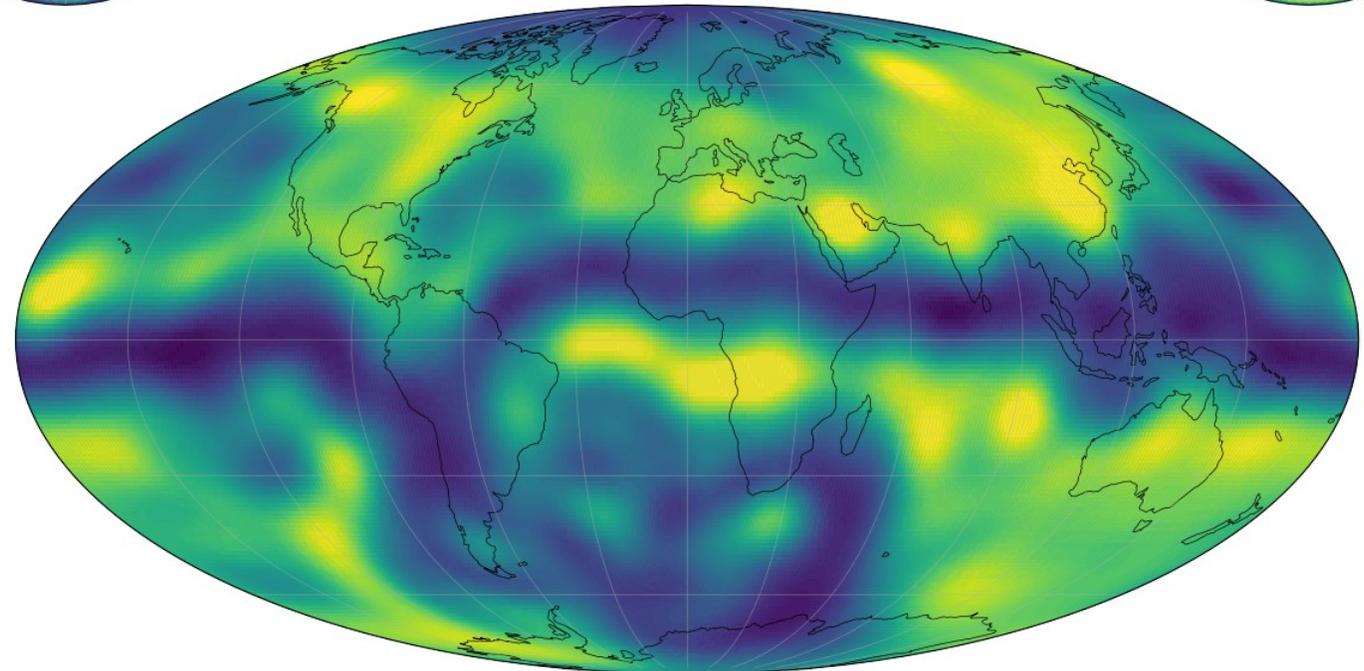
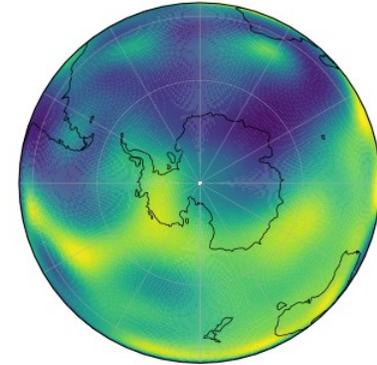
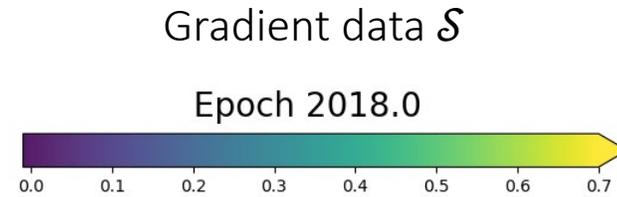
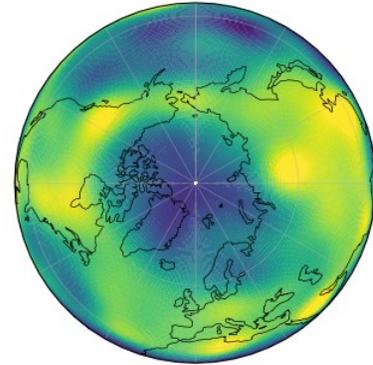
- Averaging functions (AF) are a great way to **visualise spatial resolution**
- They indicate **how well** a model estimate is **localized** at a **given point**
- We calculate AFs **across the core-mantle boundary** (and normalise values), in order to **evaluate model resolution in space**



Example of perfect averaging function at a point with max. spherical harmonic degree 14. From Whaler, Olsen, and Finlay (2016), GJI.

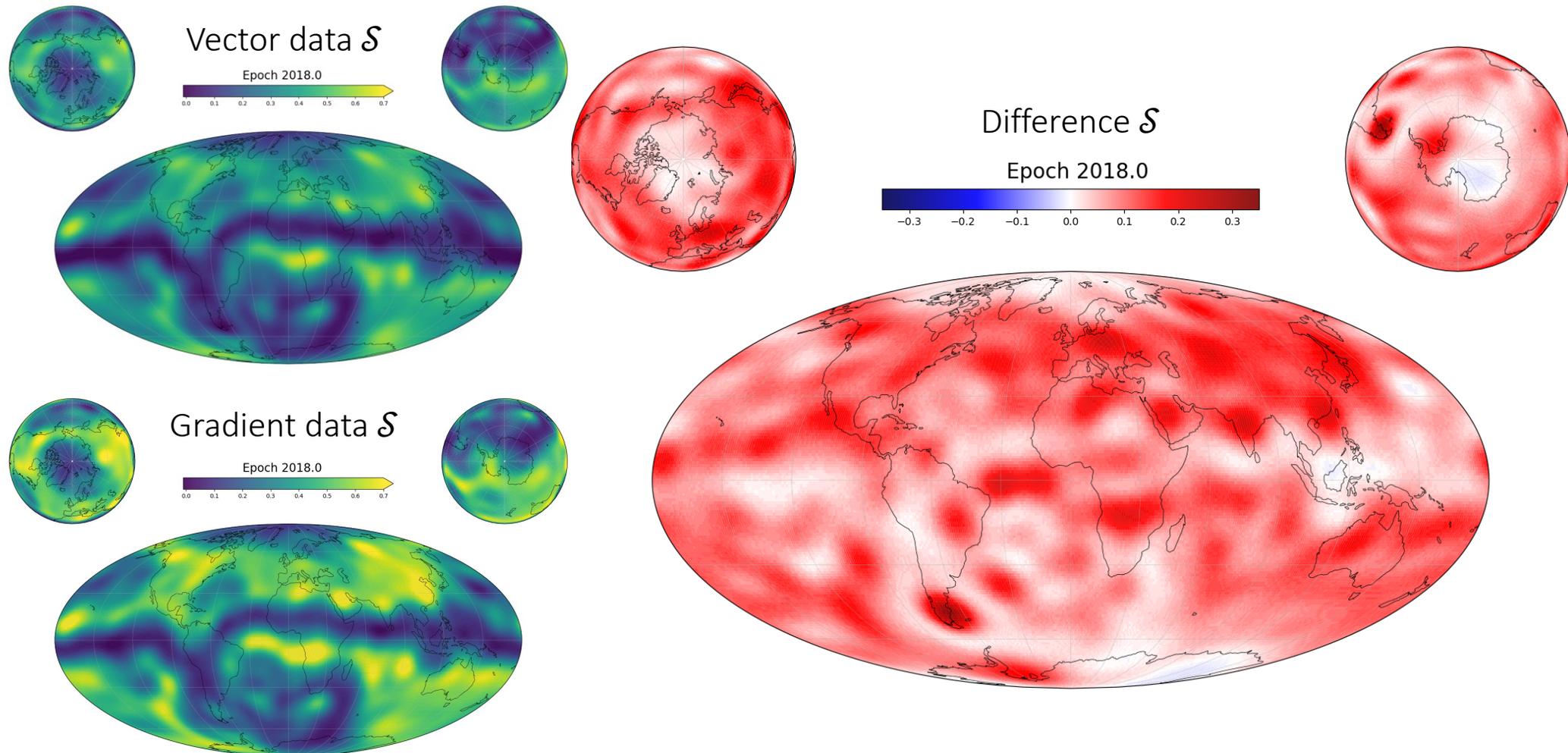
Averaging functions reveal regions of improved resolution

- Poloidal flow, \mathcal{S} , appears well resolved most locations
- Pronounced 'ambiguous' band along magnetic equator?
- Weak resolution in South Atlantic, North Pacific, and North Polar region



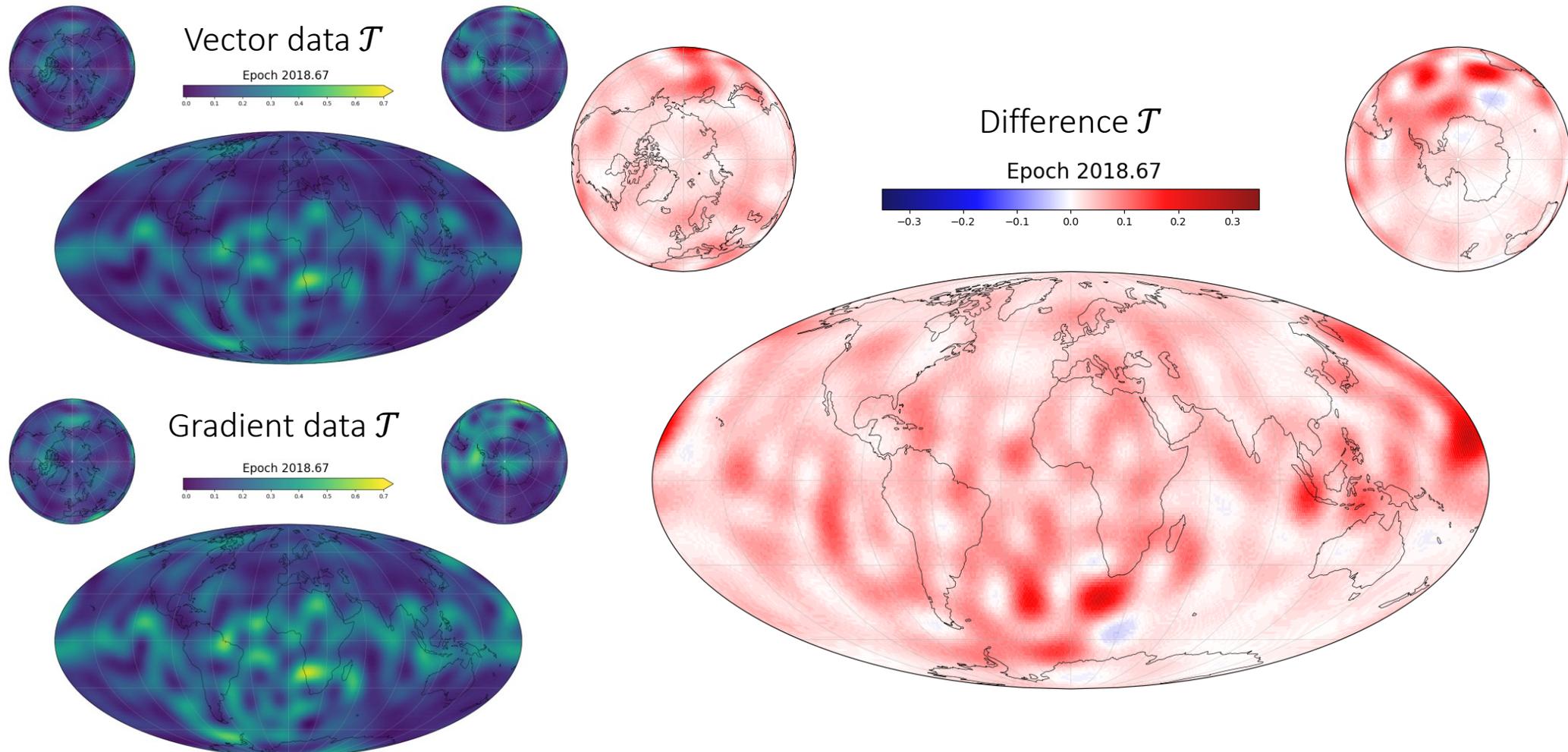
Averaging functions reveal regions of improved resolution

- Strong increase in AF (nearly) **everywhere** when using **gradients**
- Increased spatial resolution confining ambiguous patches?

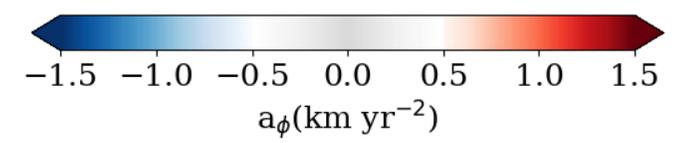
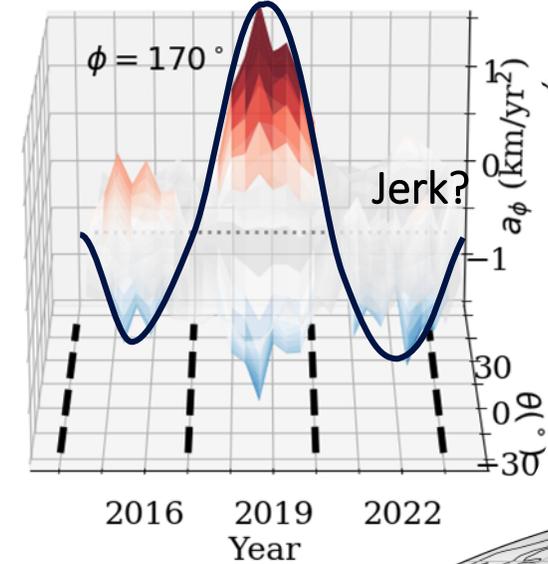
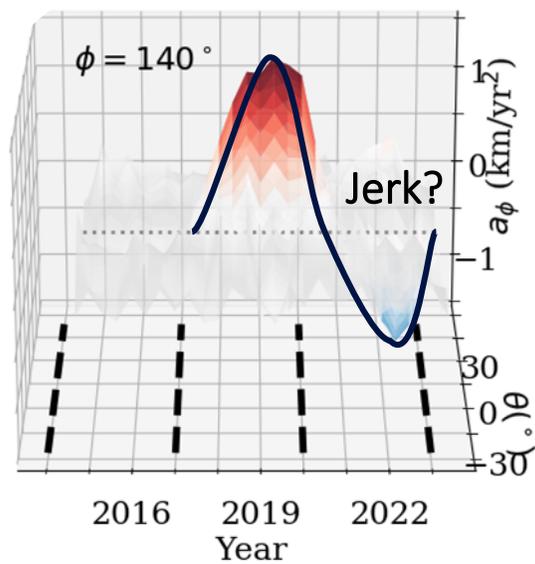


Averaging functions reveal regions of improved resolution

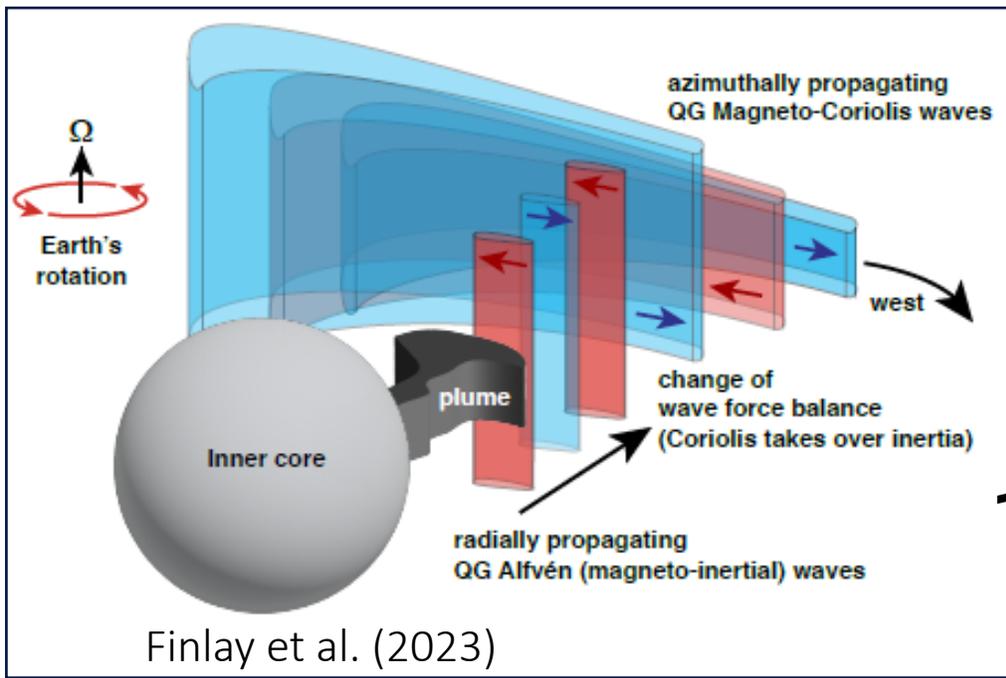
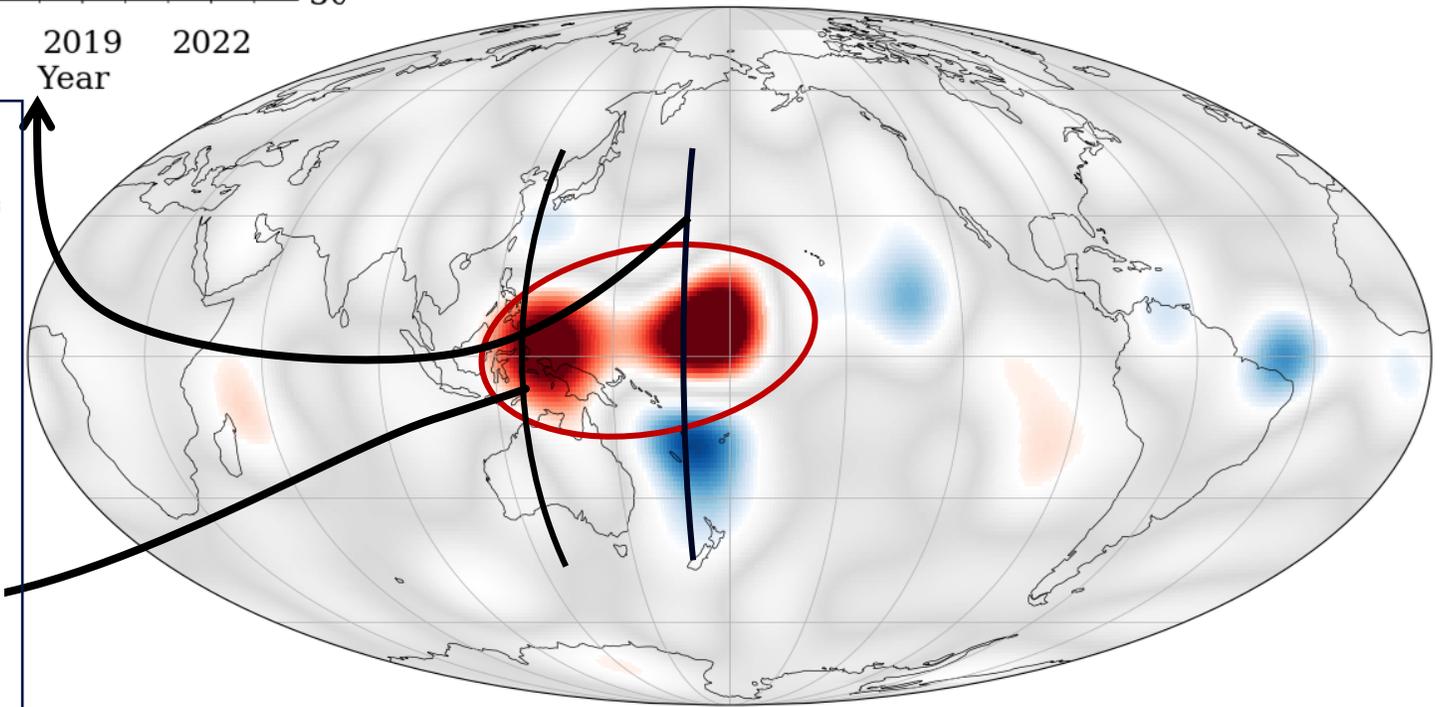
- Strong increase in AF (nearly) **everywhere** when using **gradients**
- Much weaker Afs for toroidal flow than poloidal



Result: Flow acceleration pulse in between jerks?

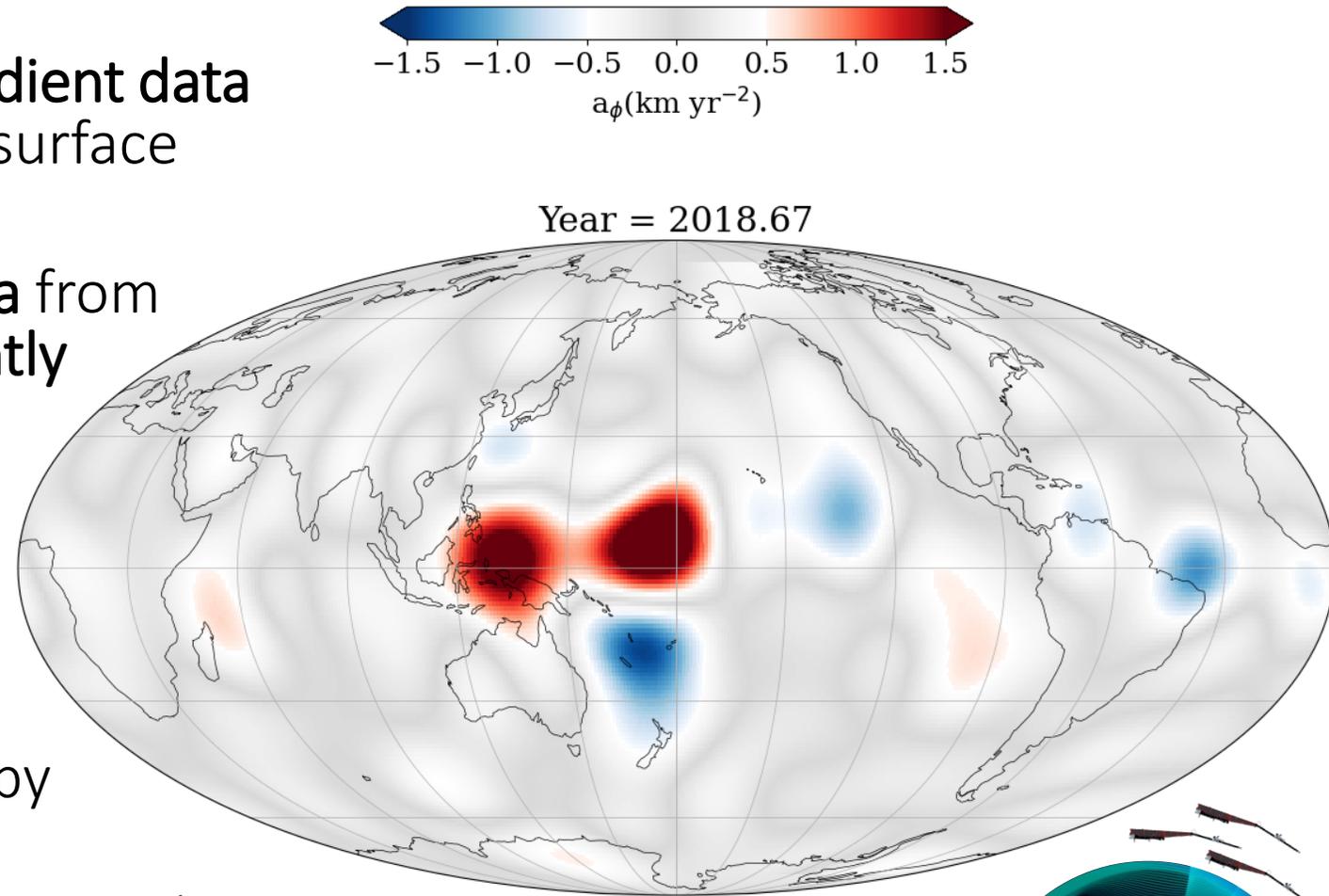


Year = 2018.67



Conclusions

- We inverted vector and spatial gradient data from CHAMP and Swarm for core-surface flow
- We found that spatial gradient data from Swarm resolved the flow significantly better than vector data
- ...whereas spatial gradient and vector data from CHAMP performed comparatively.
- Our flow models suggests that the 2017 and 2020 jerks were caused by a wave-like pulse in a_ϕ
 - This suggests a new jerk in the Pacific around 2023.



Happy
birthday to
Swarm!



Predicting variations in length-of-day (ΔLOD)

- The **minimum acceleration** models do a **poor job** at predicting ΔLOD variations
- **TO-like** models predict **better**
- **Equal performance** for **gradients** and **vectors**

