

Lessons for EarthCARE from the assimilation of cloud-affected Aeolus backscatter

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ATLID and ALADIN are on the same side!

- Previously, all cloud lidar assimilation experiments at ECMWF have used CALIPSO observations as a proxy for EarthCARE data.
 - Aeolus ALADIN lidar is technically much closer related to EarthCARE's ATLID than the CALIPSO lidar, despite its primary goal of measuring line-of-sight winds.
 - Several activities investigating the benefit of Aeolus aerosol backscatter (e.g., AEOLUS DISC), but, so far, cloud has mainly been seen as noise that must be removed.
- Can Aeolus observations be used to prepare for ATLID assimilation?
 - How useful is Aeolus cloud backscatter for NWP?
 - Can Rayleigh backscatter be assimilated?

Vertical resolution and pointing angle are the major differences between ATLID and ALADIN

Characteristic	EarthCARE ATLID	Aeolus ALADIN
Specification	3° off-Nadir 355 nm 0.62 m diameter telescope 35 µrad receiver field of view	37.5° off-Nadir 355 nm 1.50 m diameter telescope 19 µrad receiver field of view
Altitude	400 km	400 km
Resolution	103 m vertical 0.285 km horizontal 0.03 km footprint	500 m - 1000 m vertical ~5 km horizontal 0.009 km footprint
Misc.	High-spectral-resolution receiver with Rayleigh and Mie co-polar and total cross-polar channels	High-spectral-resolution receiver with Rayleigh and Mie copolar channels only

...requires some adaptations to observation operator

Observation-side

Data products and pre-processing



Initially L2A total attenuated backscatter. Capability for individual HSRL channels



Mie attenuated backscatter

Rayleigh attenuated backscatter

Observation screening

Comprehensive, based upon CALIPSO studies and L2A cloud mask.

Basic screening on FG departures and minimum backscatter values

Observation errors

Error inventory approach, combining representativity error*, instrument error and forward model error

Fixed value of 10 dB

*see Fielding and Stiller 2019, JGR

Model-side

Microphysical assumptions



PSD: Field et al., 2007

Single scattering properties: Baran?? 2010 mixture with fixed lidar ratio (consistent with CAPTIVATE retrievals)

JOINT

ALADIN

Observation operator

Triple column approach accounting for sub-grid condensate variability and cloud overlap. Platt approximation for multiple scattering. + adaptations for ALADIN

See Fielding and Janisková 2020, QJRMS

Bias correction

Initially none, then climatological based upon temperature and region.

None

4D-Var

Adapting the observation operator for ALADIN

Apparent backscatter

Platt coefficient

Optical depth due to clouds

Optical depth due to gases

$$\beta'(r) = (1 - \delta)\beta(r)e^{-\frac{2(\eta\tau_{cloud} + \tau_{gas})}{\cos\theta}}$$

Unattenuated backscatter

Slant angle

Depolarization ratio

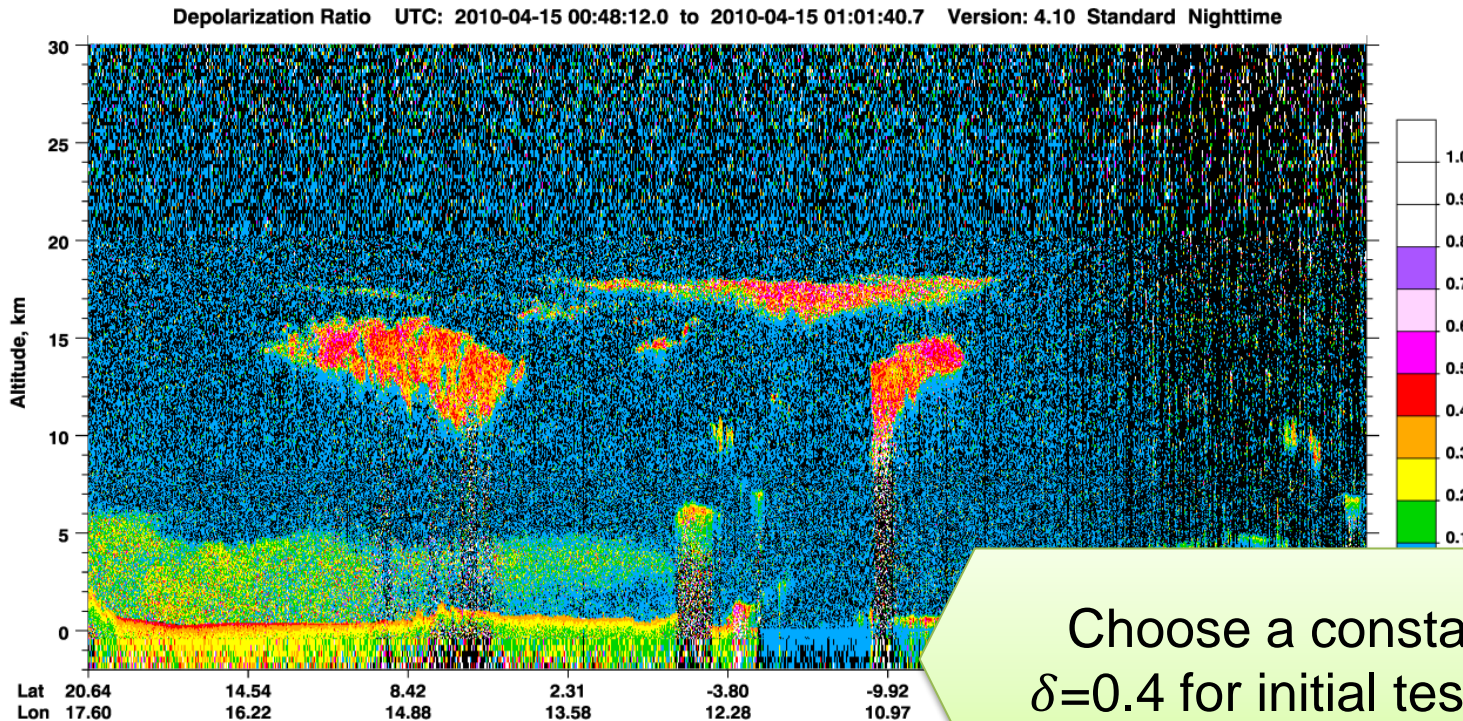
The diagram illustrates the equation for apparent backscatter, $\beta'(r)$, which is a function of the unattenuated backscatter $\beta(r)$, the depolarization ratio δ , and the slant angle θ . The equation is $\beta'(r) = (1 - \delta)\beta(r)e^{-\frac{2(\eta\tau_{cloud} + \tau_{gas})}{\cos\theta}}$. The term $\eta\tau_{cloud}$ represents the optical depth due to clouds, and τ_{gas} represents the optical depth due to gases. The term $\cos\theta$ is the slant angle. The term $(1 - \delta)$ is the depolarization ratio. The term $\beta(r)$ is the unattenuated backscatter. The term η is the Platt coefficient.

Modifications for ALADIN

Determining the depolarization ratio

- Unlike spherical particles, such as water droplets, ice particles are generally strongly depolarizing.
- Depolarization ratio can be measured empirically or computed numerically from assumed ice particle shapes.

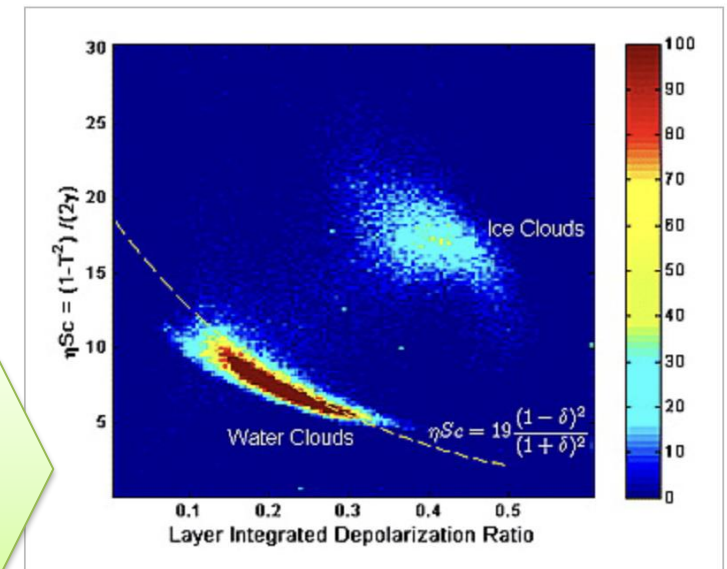
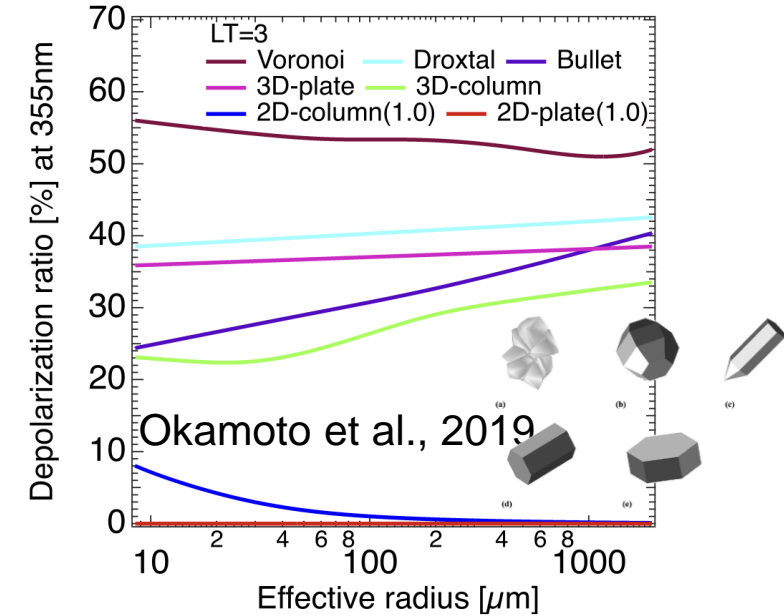
Observed using CALIPSO (532 nm)



Choose a constant $\delta=0.4$ for initial testing

Courtesy NASA (CALIPSO user guide)

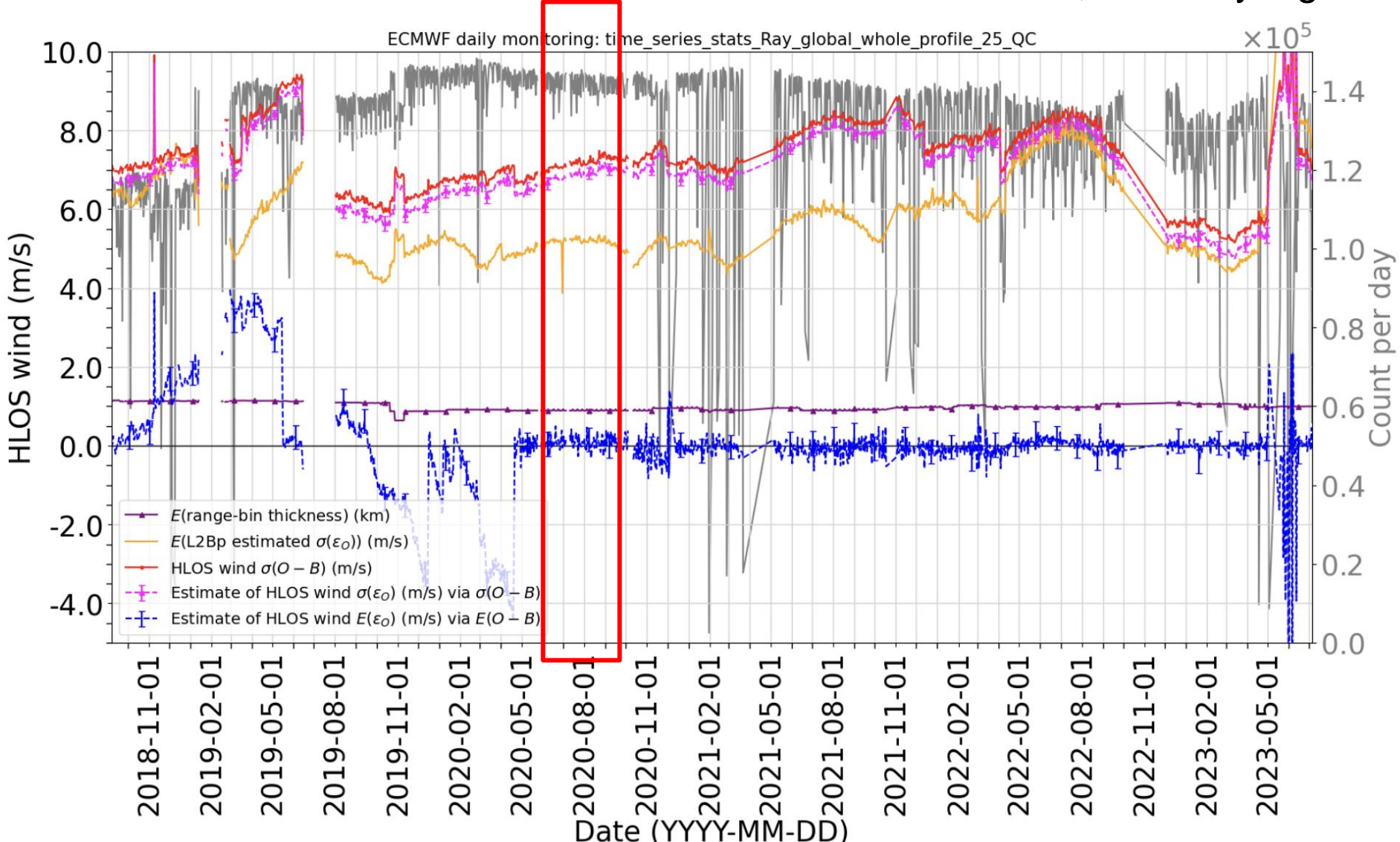
Theoretical (355 nm)



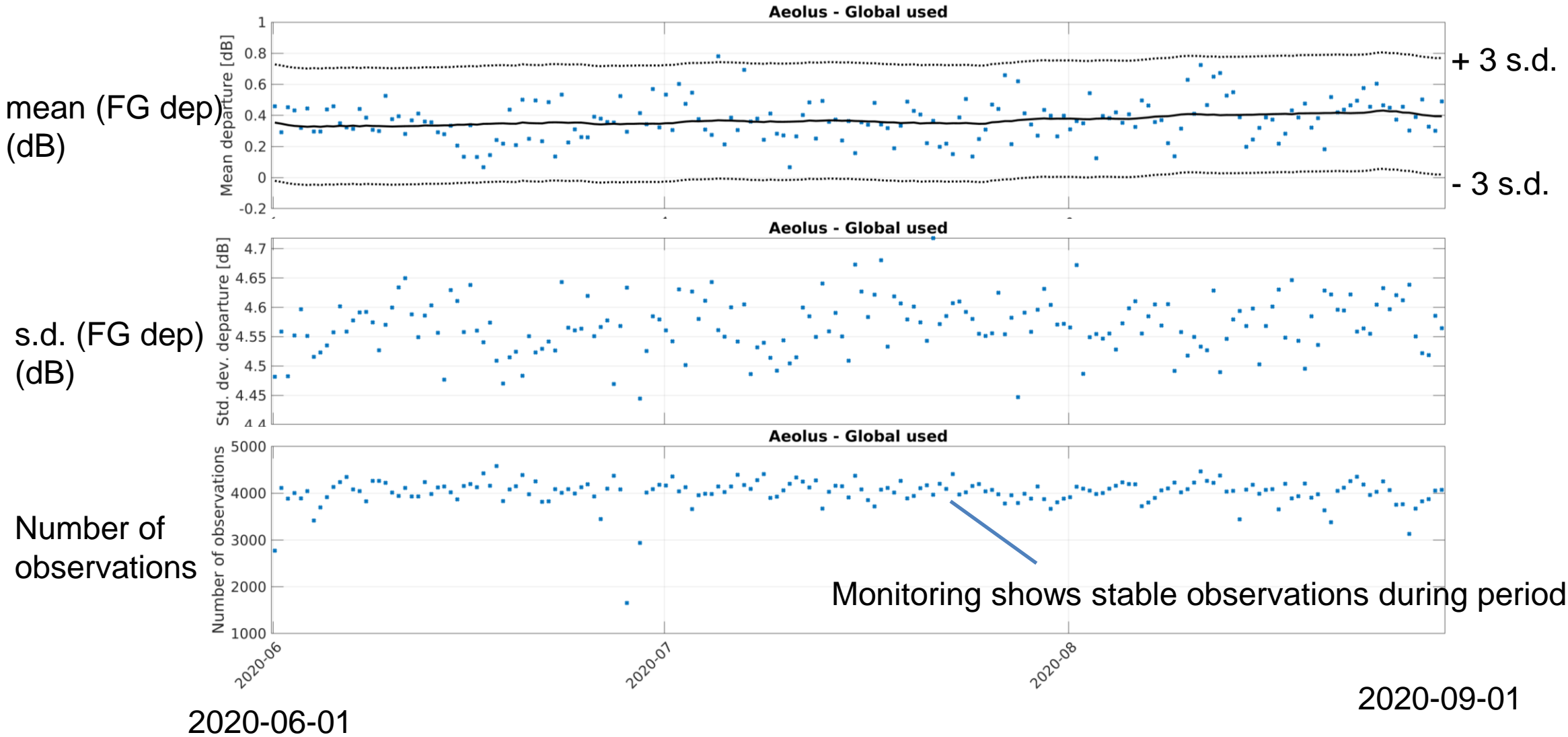
Hu, 2007

Choosing test period --- ECMWF monitoring of Aeolus HLOS winds

Mid-mission: lower errors, relatively high laser energy

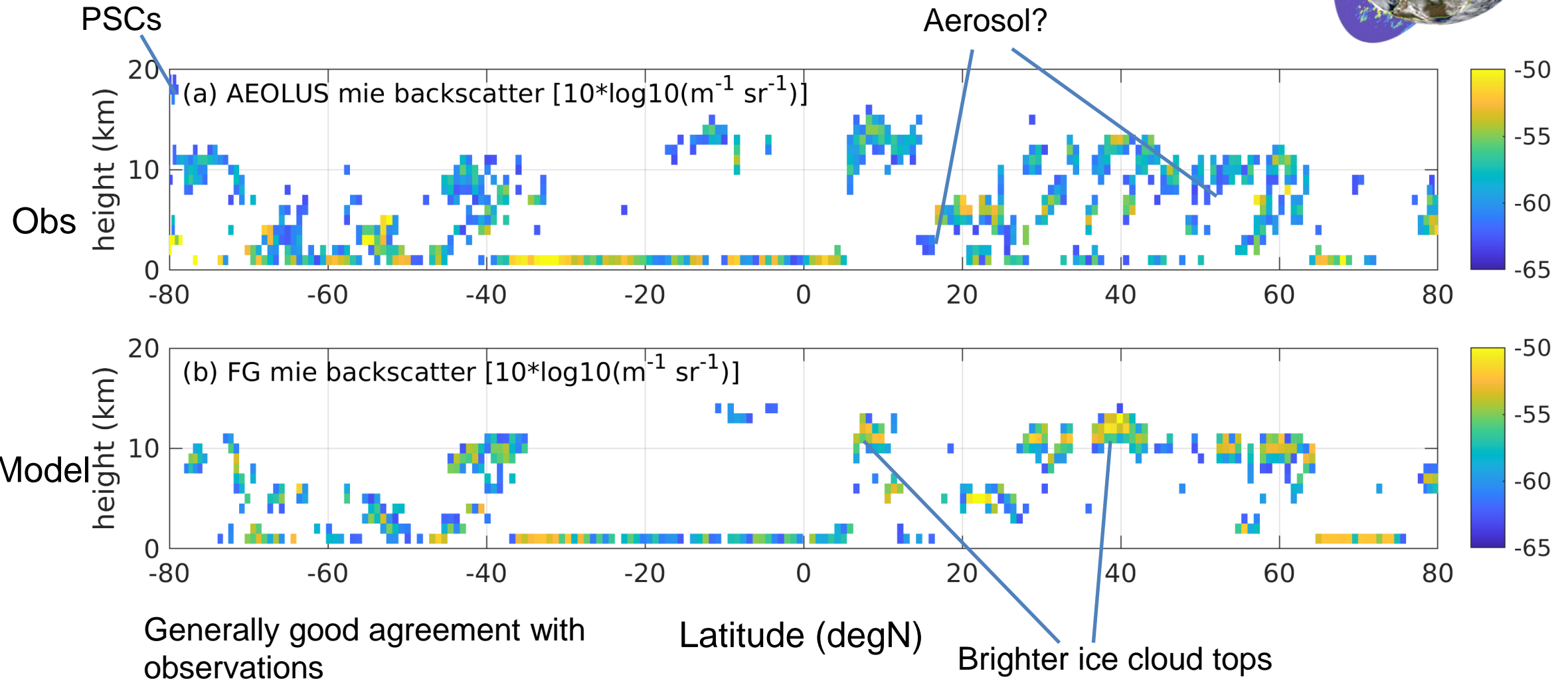
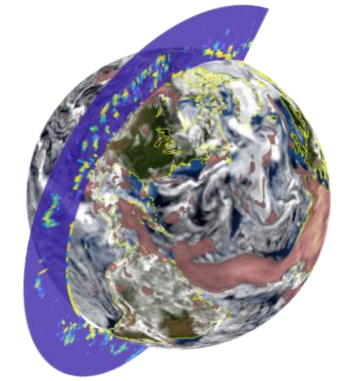


12-hourly monitoring of Aeolus L2A Mie attenuated backscatter



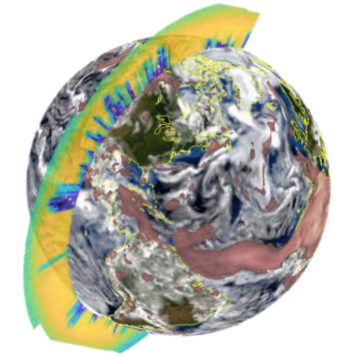
Aeolus orbit– 1st June 2020

Mie attenuated backscatter



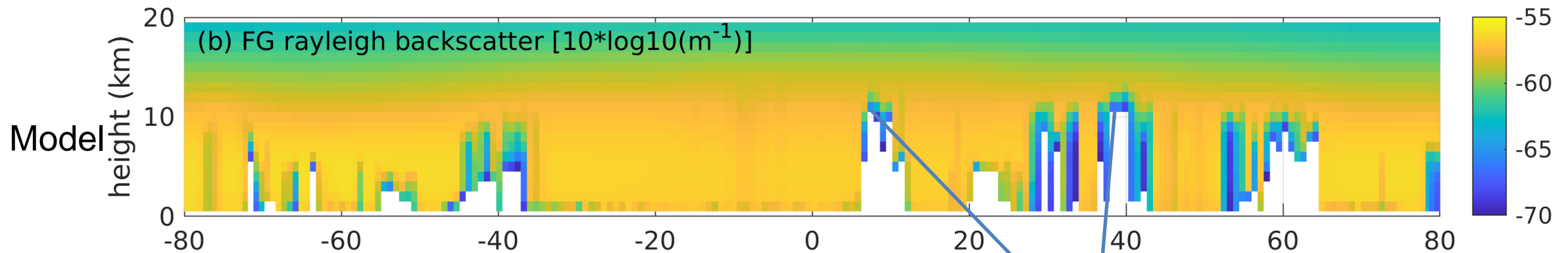
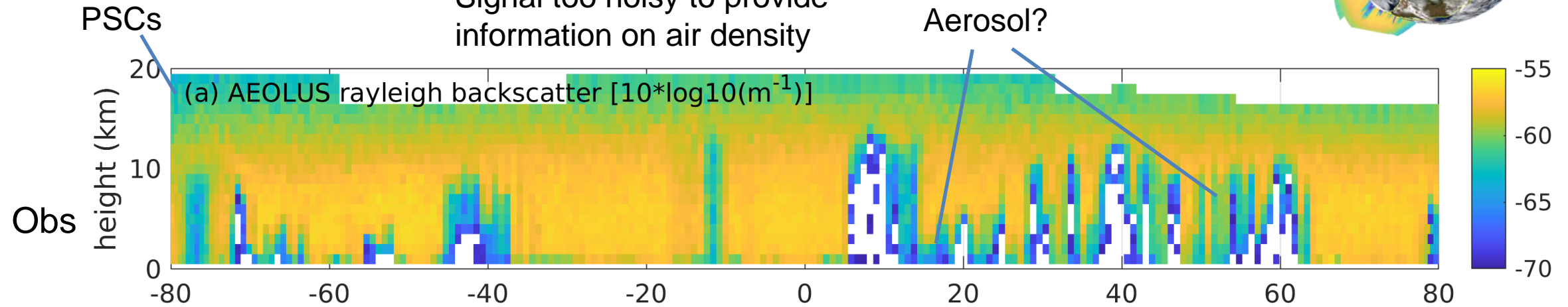
Aeolus orbit– 1st June 2020

Rayleigh attenuated backscatter



Signal too noisy to provide information on air density

Aerosol?

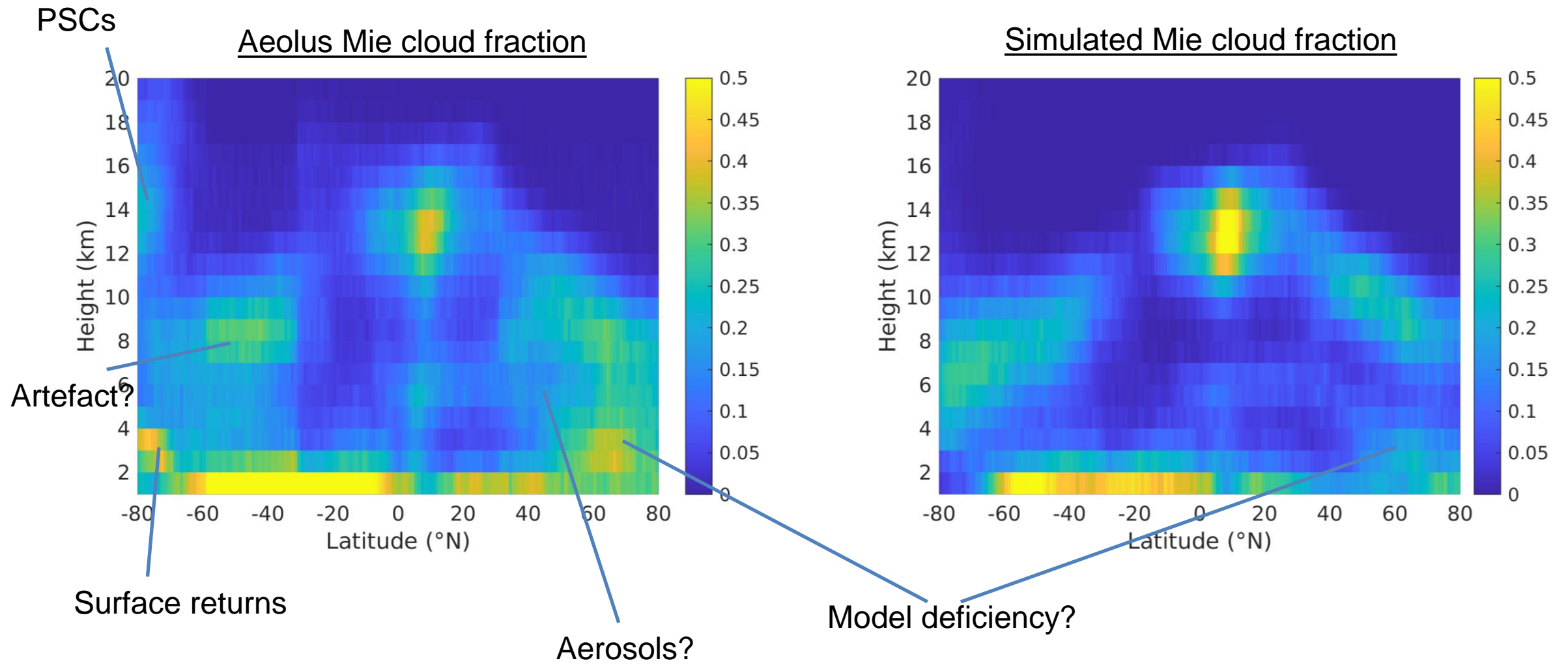


Generally good agreement with observations

Latitude (degN)

Rayleigh signal not affected by depolarization

Climatology of observed and simulated cloud fraction – June 2020

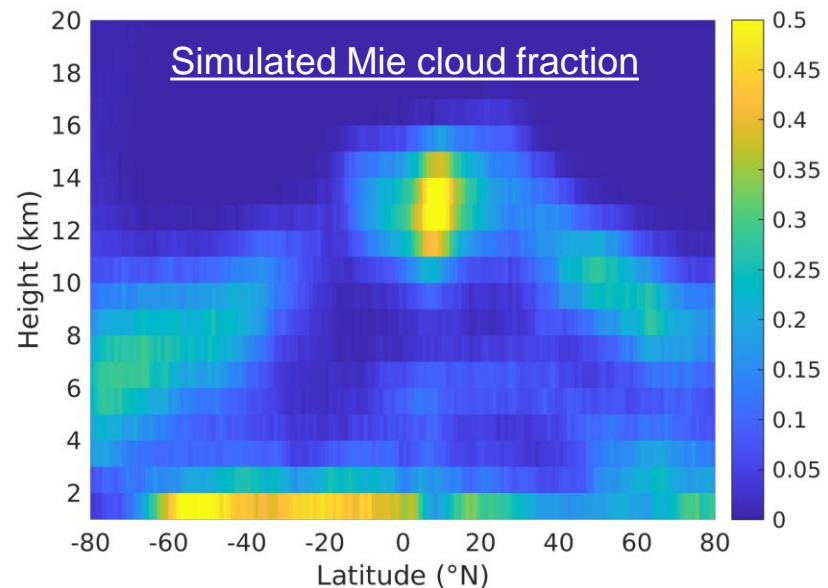
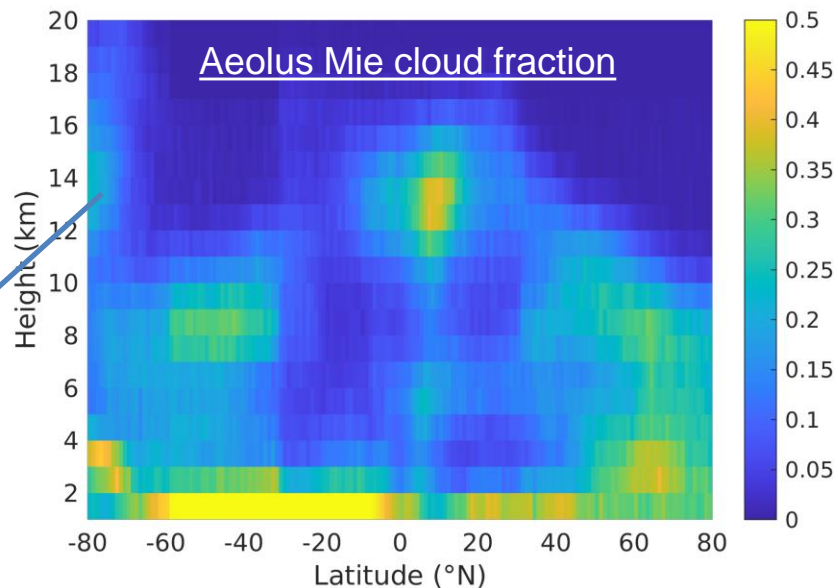


Comparison of Aeolus and CALIPSO observed cloud amount

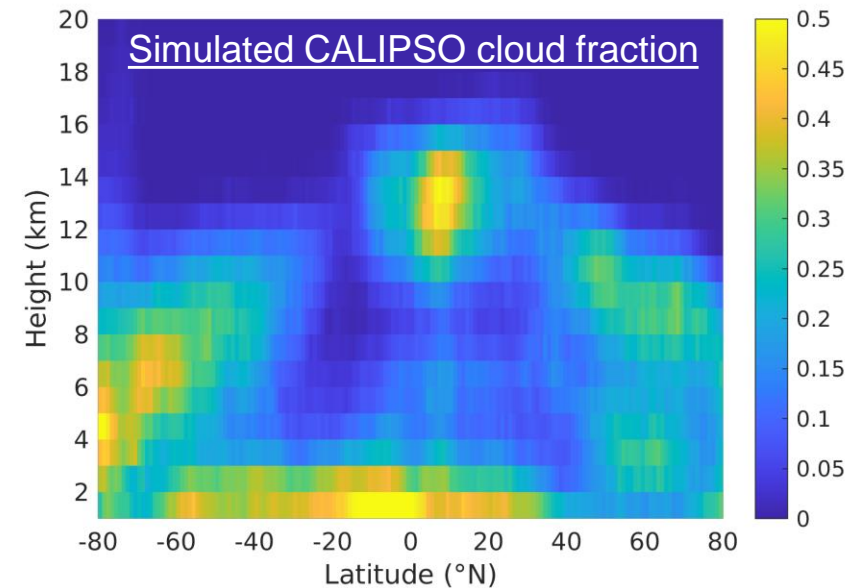
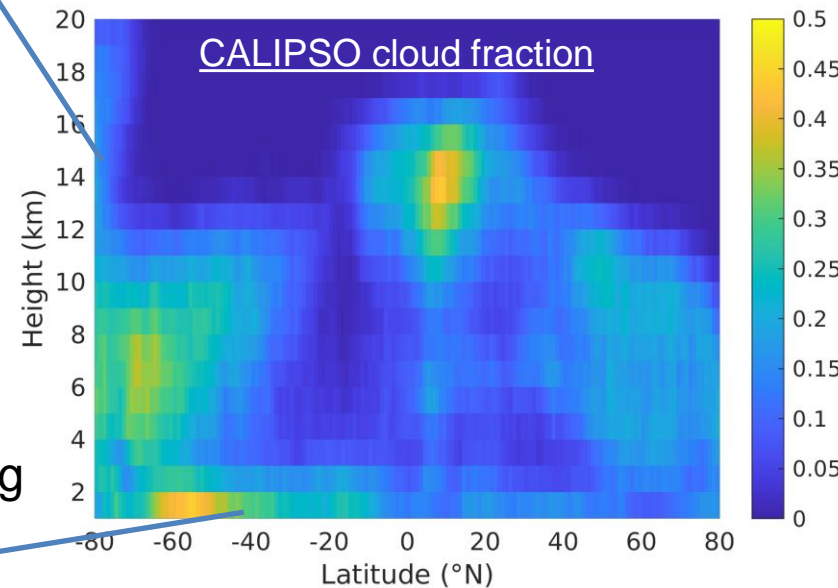
More noise in Aeolus backscatter?

PSCs brighter at 355 nm

More screening in CALIPSO obs?



June 2020



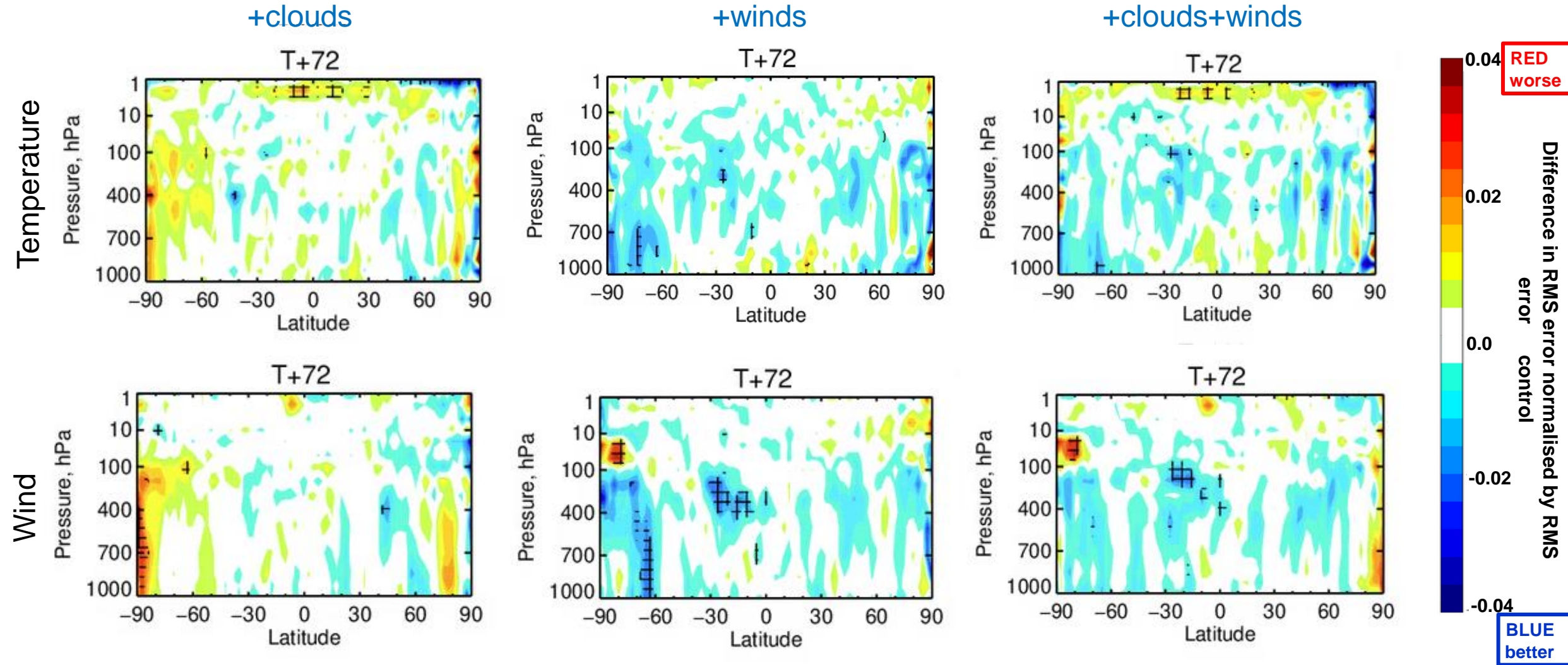
June 2010

See Feofilov et al., 2022 AMT for more detailed analysis

4D-Var experiment setup

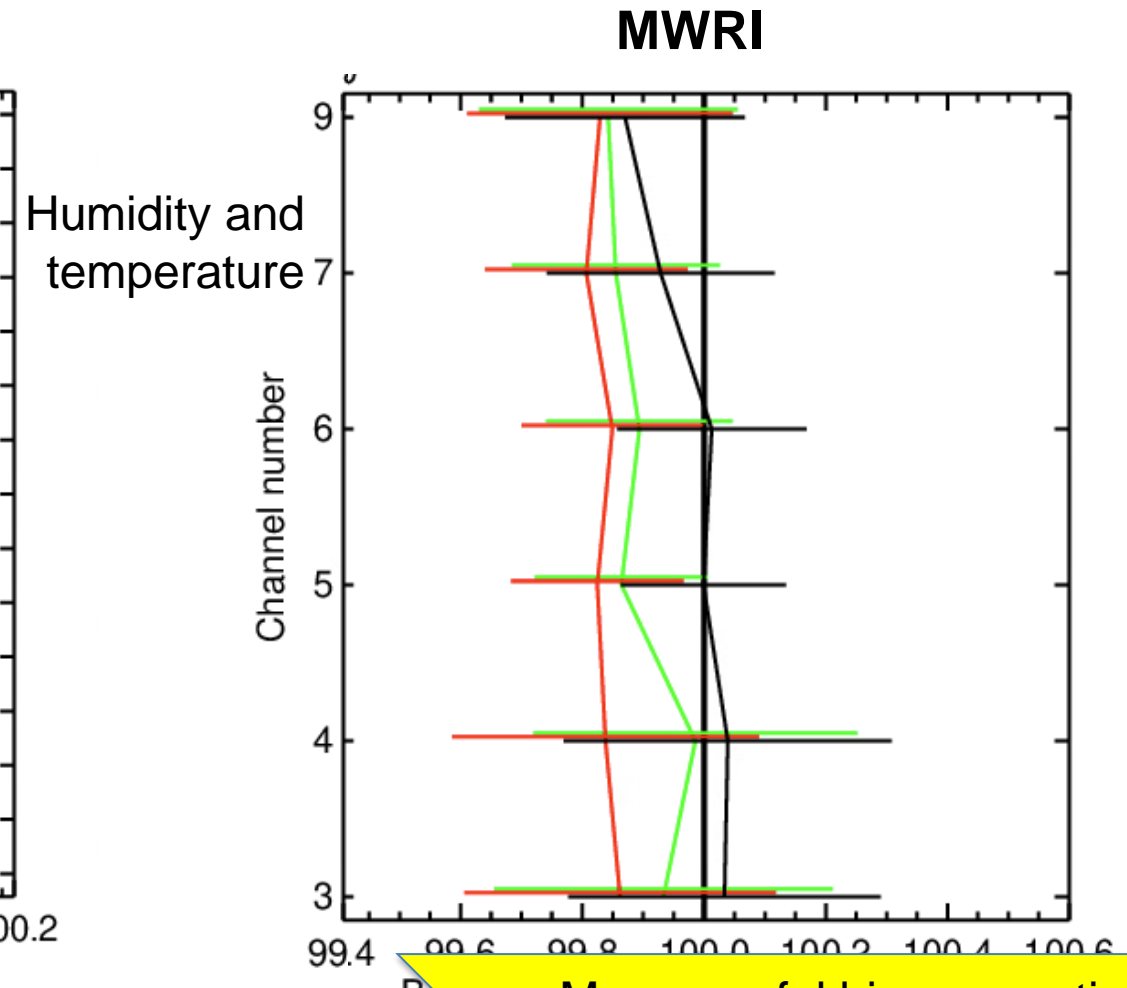
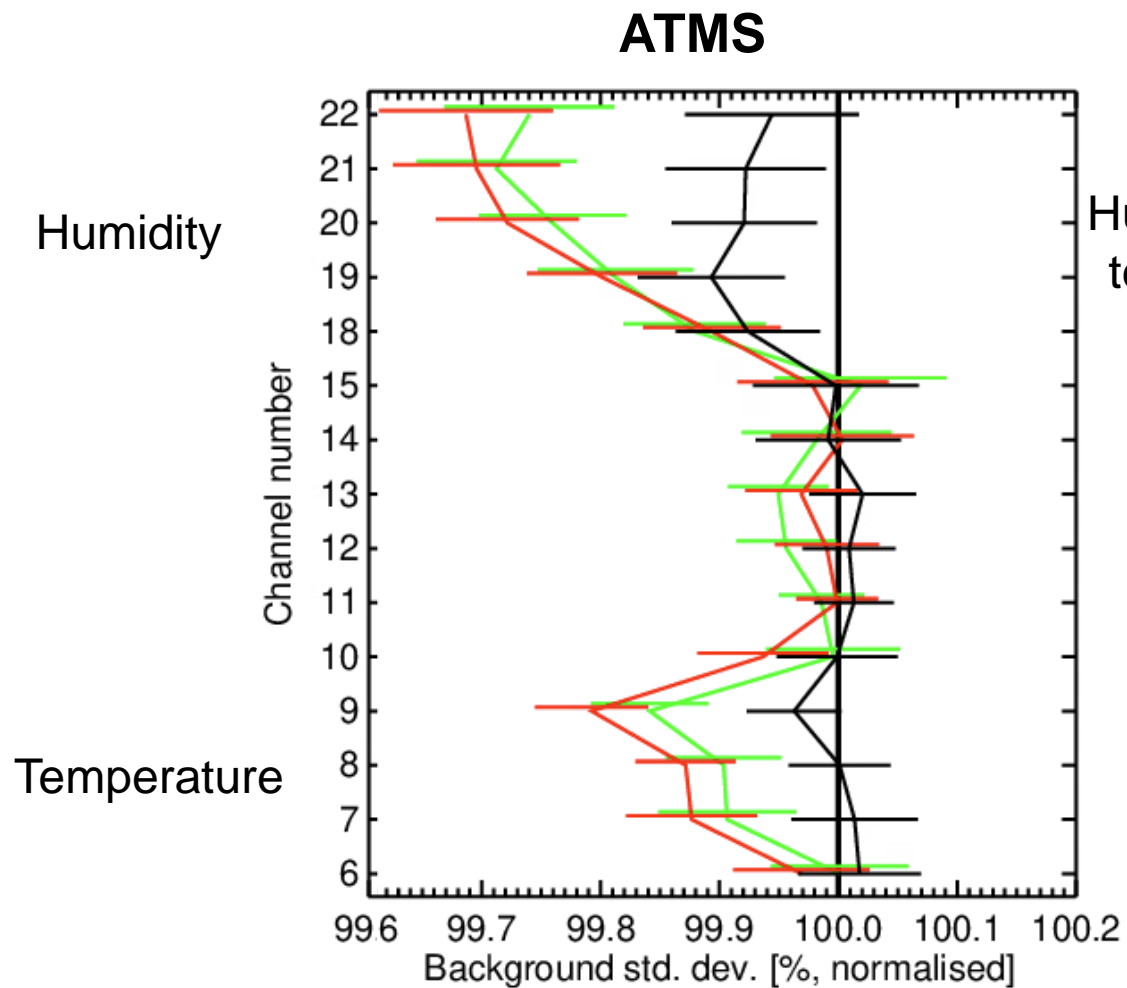
- CY48R1 4D-Var experimentation using a horizontal resolution of TCo639 spectral truncation (corresponding to ~ 18 km on a cubic octahedral grid) and 137 levels:
 - 3-month period: 1 June 2020 – 31 August 2020
- Measurements of L2A Mie and Rayleigh attenuated backscatter (at 355 nm, ALADIN) superobbed to (O160-> ~ 72 km) and 1 km vertical heights.
- Performed experiments:
 - Control – reference run, run with all regularly assimilated observations except Aeolus winds
 - **+winds** – operational configuration including Aeolus HLOS winds – see Rennie et al., 2021
 - **+clouds** – experimental run including attenuated Mie and Rayleigh attenuated backscatter lidar observations on top of other regularly assimilated observations
 - **+winds+clouds** - experimental runs including attenuated Mie and Rayleigh attenuated backscatter lidar observations and HLOS winds

Assimilating Aeolus cloud obs. broadly neutral on forecasts of large-scale variables



3-month period: 1 June 2020 – 31 August 2020

Assimilating Aeolus cloud Mie and Rayleigh backscatter has slight positive benefit on microwave all-sky obs!



- clouds only
- +mie+rayleigh+winds
- winds only
- 100% = 48r1 control (no aeolus)

More careful bias correction and observation error tuning should increase impact

Summary

- Aeolus cloud backscatter observations have been assimilated into a global model for the first time!
- Initial results show impact is broadly neutral, but slight additional positive benefit when simultaneously assimilating winds and cloud information.
- Cloud observations from Aeolus could be useful for model evaluation – Aeolus sees clouds at different point in diurnal cycle. PSCs are clearly visible!
- More careful cloud/aerosol screening required for comparison with CALIPSO
 - Can Aeolus observations be used to prepare for ATLID assimilation? YES
 - How useful is Aeolus cloud backscatter for NWP? YES
 - Can Rayleigh backscatter be assimilated? YES (but more analysis required)

